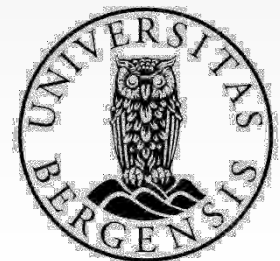
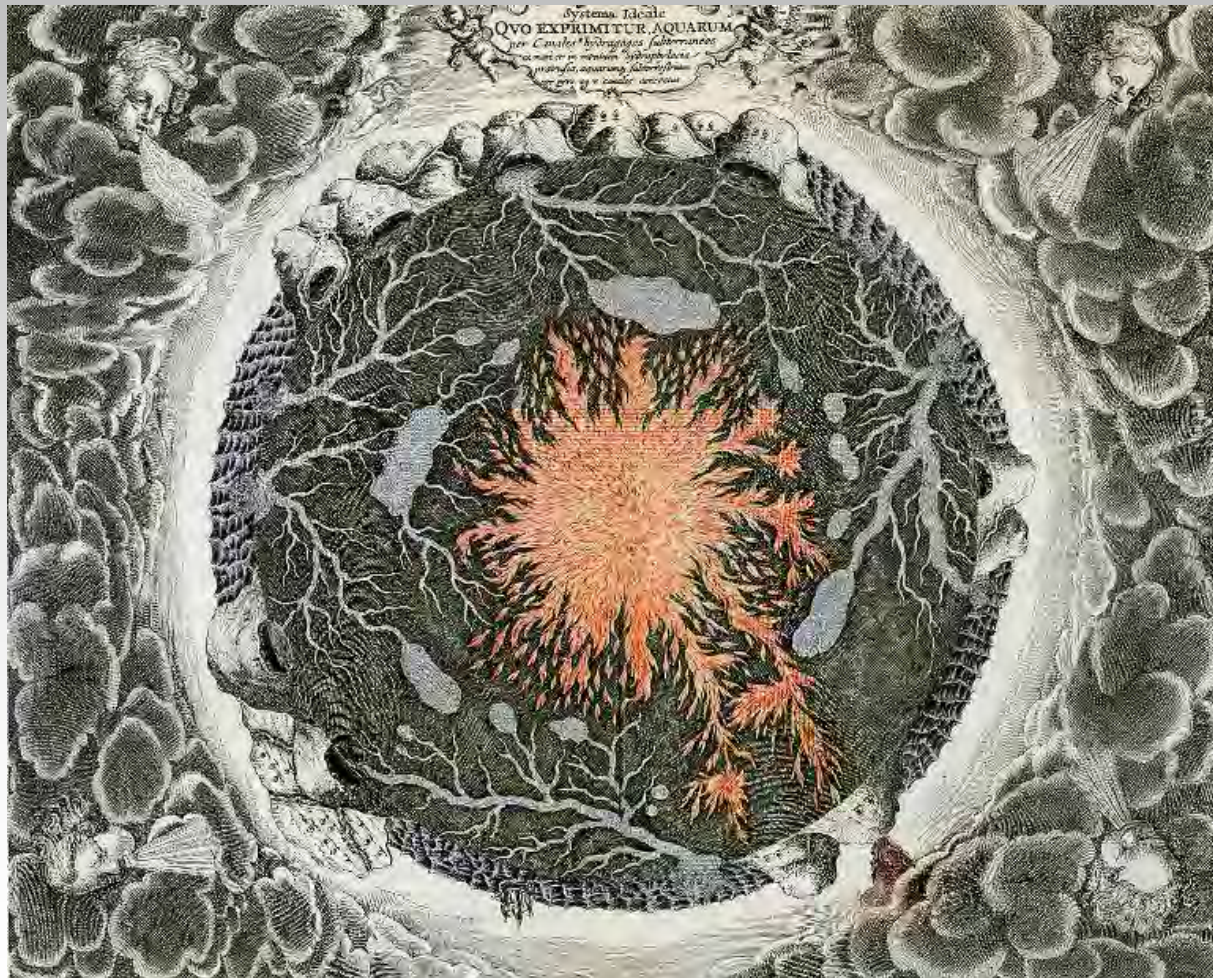


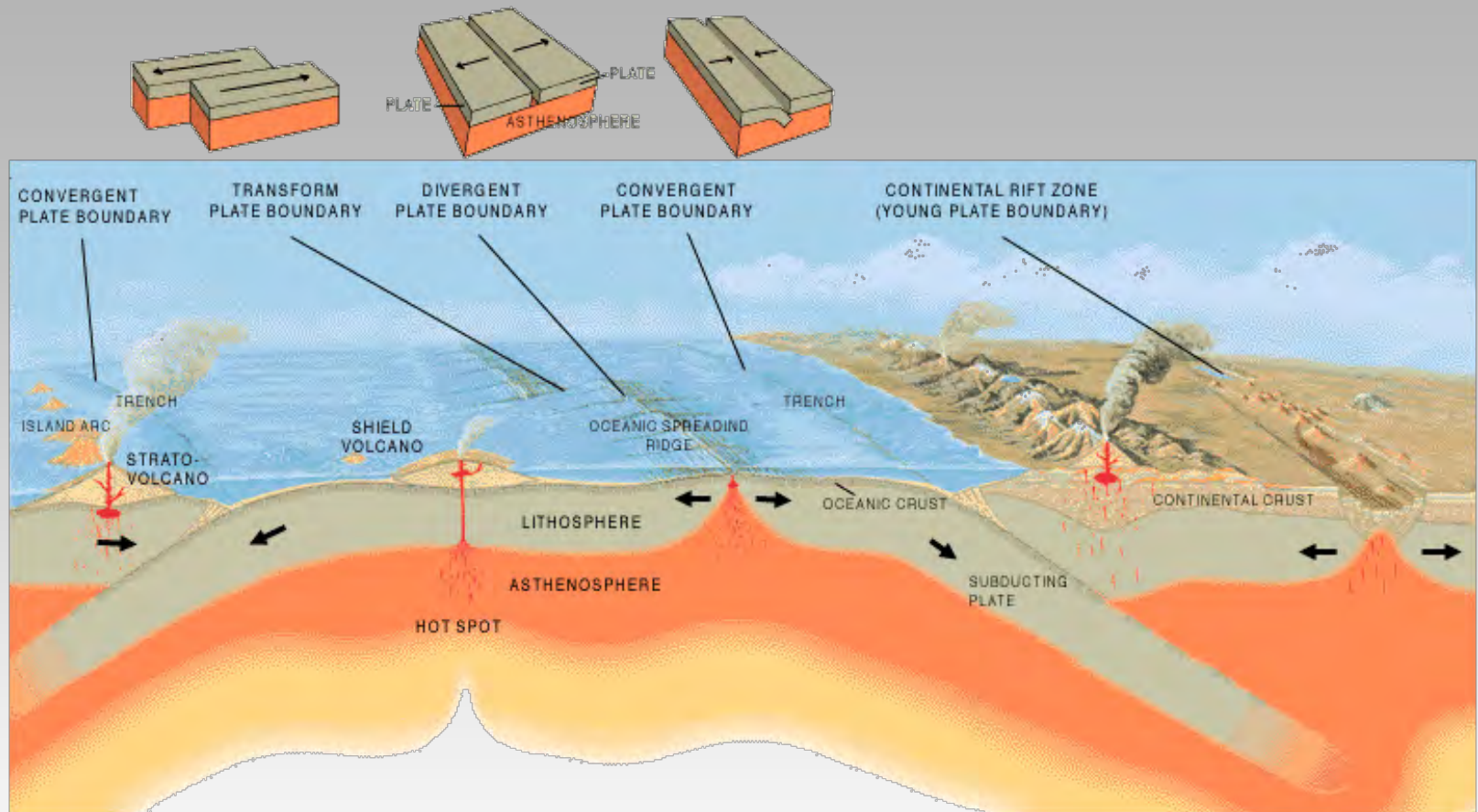
Mapping the distribution of fluids in the crust and lithospheric mantle utilizing geophysical methods

Martyn Unsworth, University of Alberta, CANADA

Stéphane Rondenay, University of Bergen, NORWAY



13.1 : Introduction



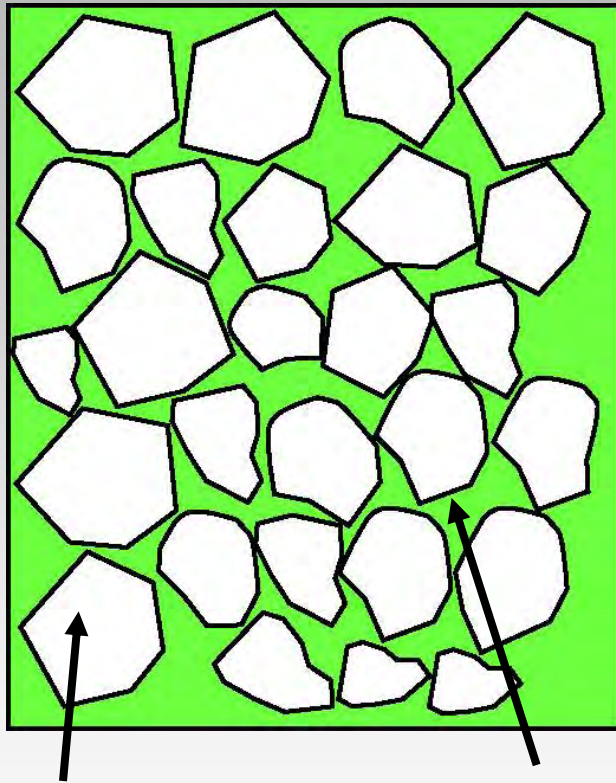
Fluid distribution at depth is important

- changes composition of crust and mantle through metasomatism
- controls rheology, mode of deformation

13.1 : Introduction - effect of fluids on rock properties

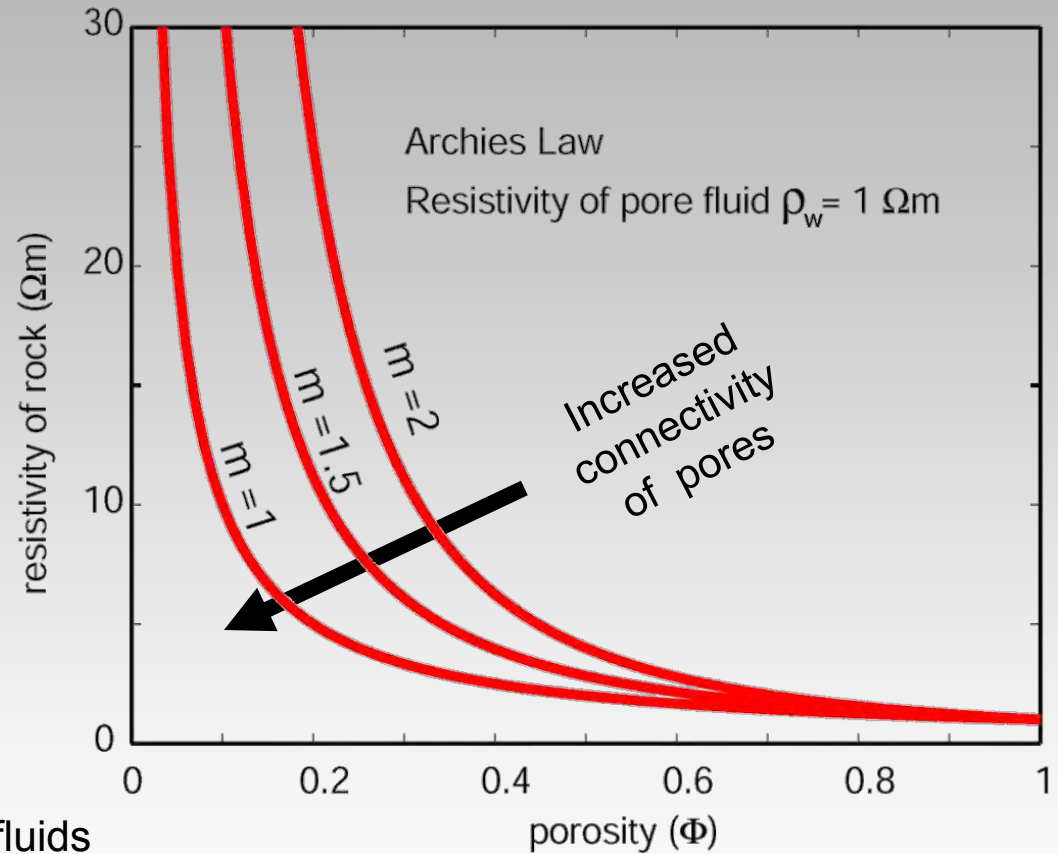
	Resistivity	Seismic properties
Free water	Decreases Sensitive to salinity	Lowers velocity Changes Poisson's ratio
Hydrous minerals Source of fluids Result of metasomatism	Decreases?	Lower velocity Increased anisotropy
Anhydrous minerals H ⁺ in olivine	Decreases Anisotropy	Lower velocity Increased anisotropy
Partial melt	Decrease	Lower velocity Attenuation Anisotropy if deformed

13.2 : Electromagnetic methods – electrical resistivity of rocks – free water



rock grains
high resistivity

Melt or aqueous fluids
low resistivity



$$\rho_o = \rho_w \Phi^{-m}$$

13.2 : Electromagnetic methods – resistivity of aqueous fluids – free water

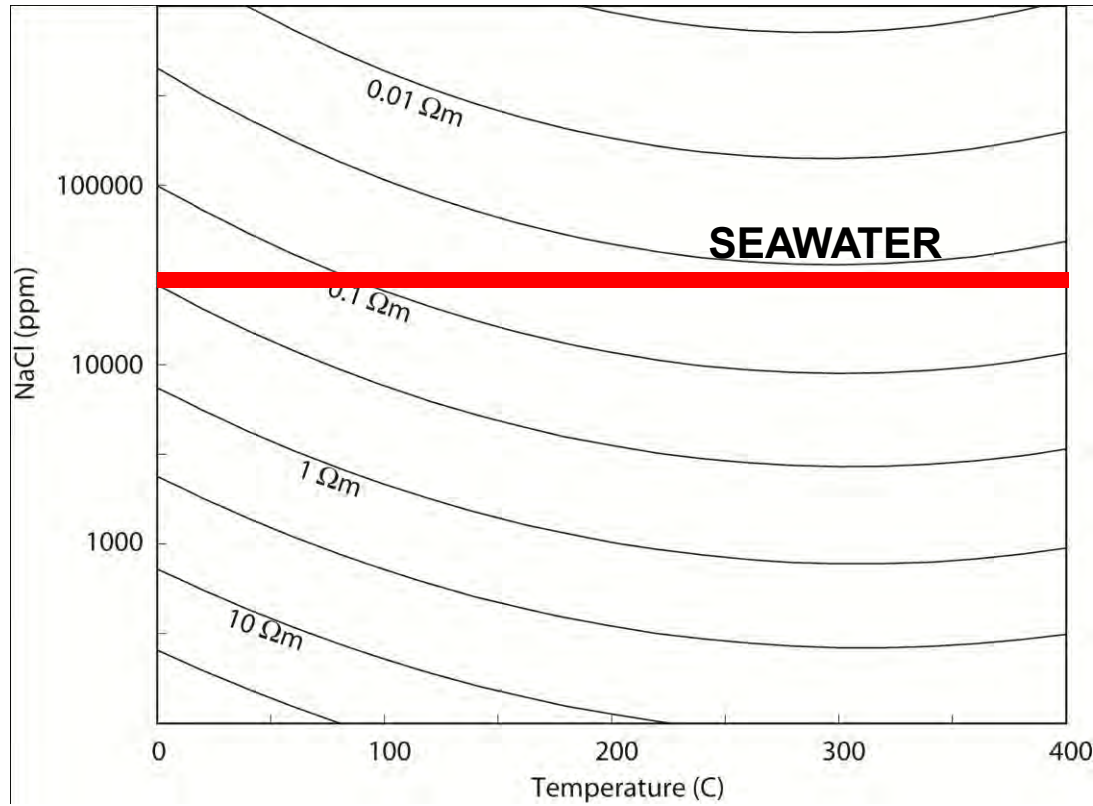


Figure 13.1

13.2 : Electromagnetic methods – resistivity of aqueous fluids – free water

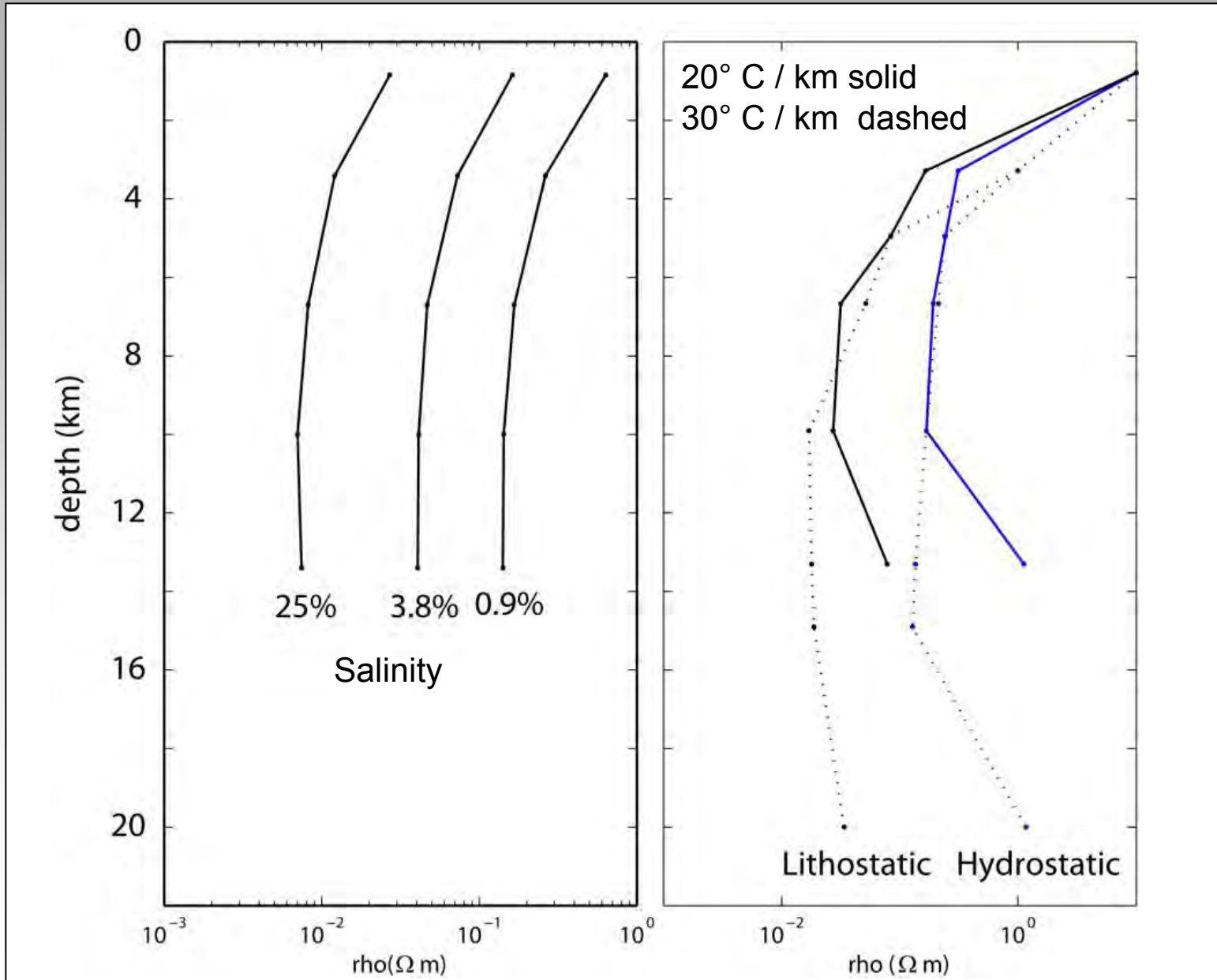


Figure 13.3
based on
Nesbitt (1993)

13.2 : Electromagnetic methods – dihedral angles – free water

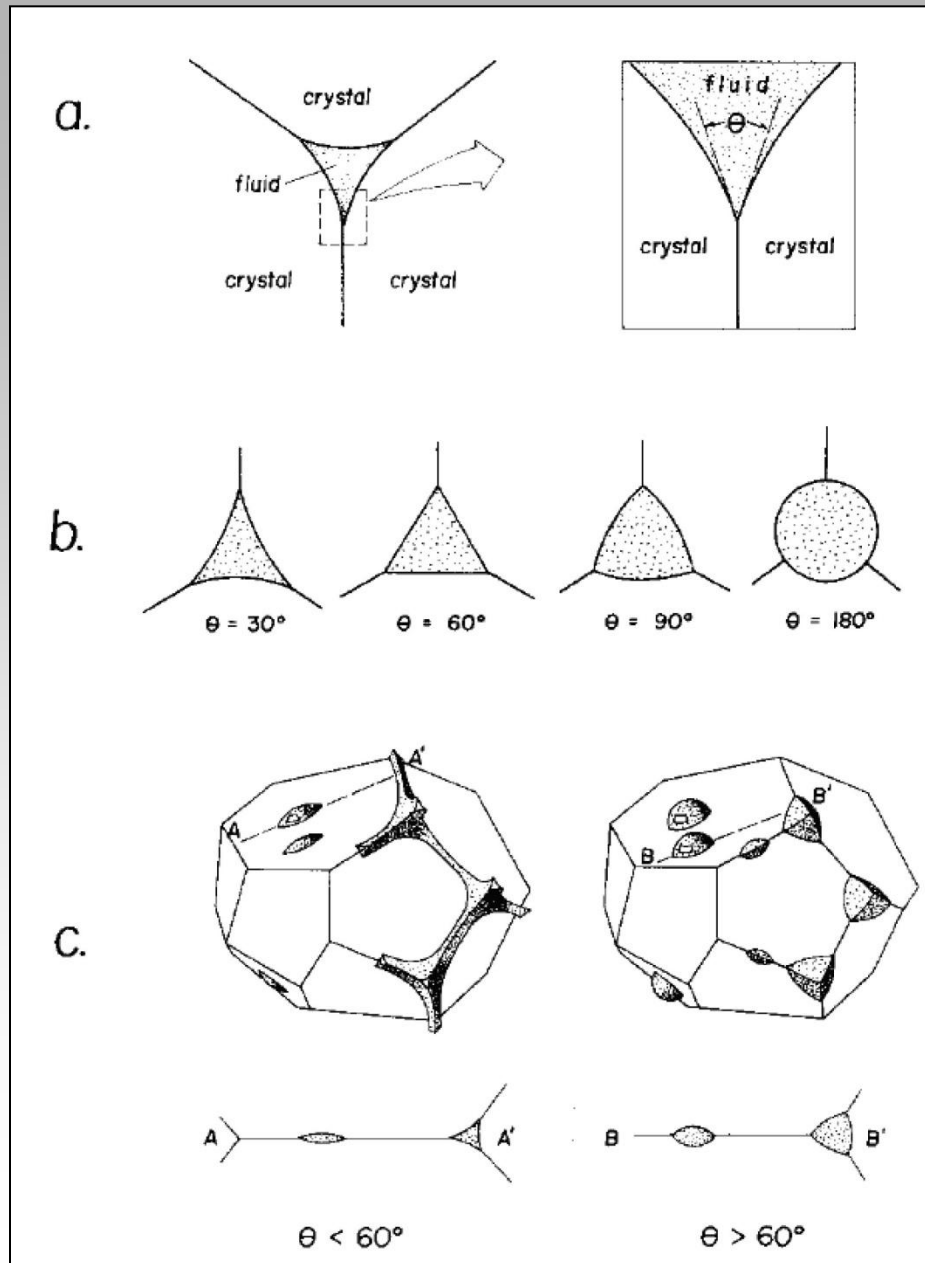
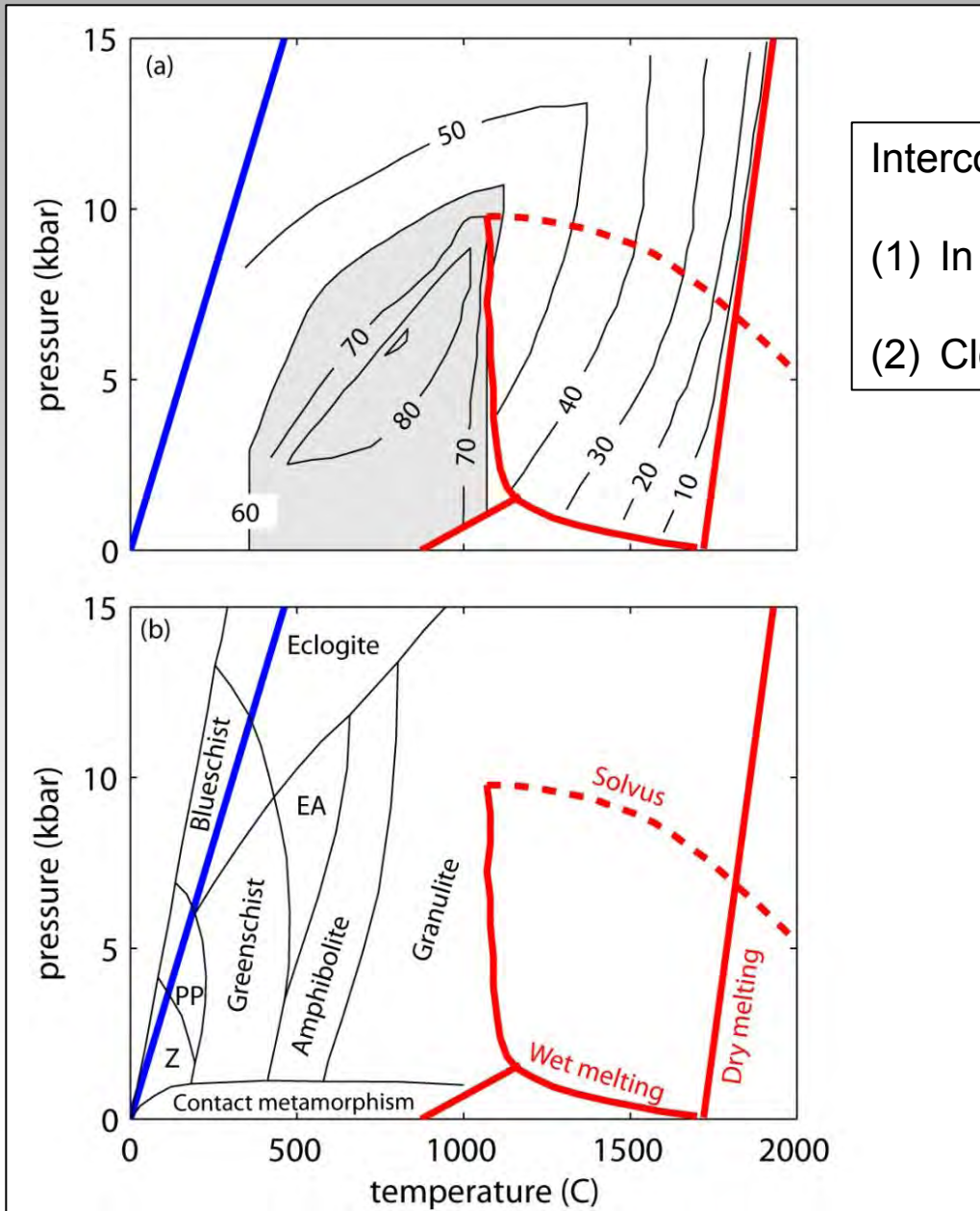


Figure 13.4

13.2 : Electromagnetic methods – dihedral angles - free water



Dihedral angle from Holness (1993)

Interconnection of aqueous fluids may occur :

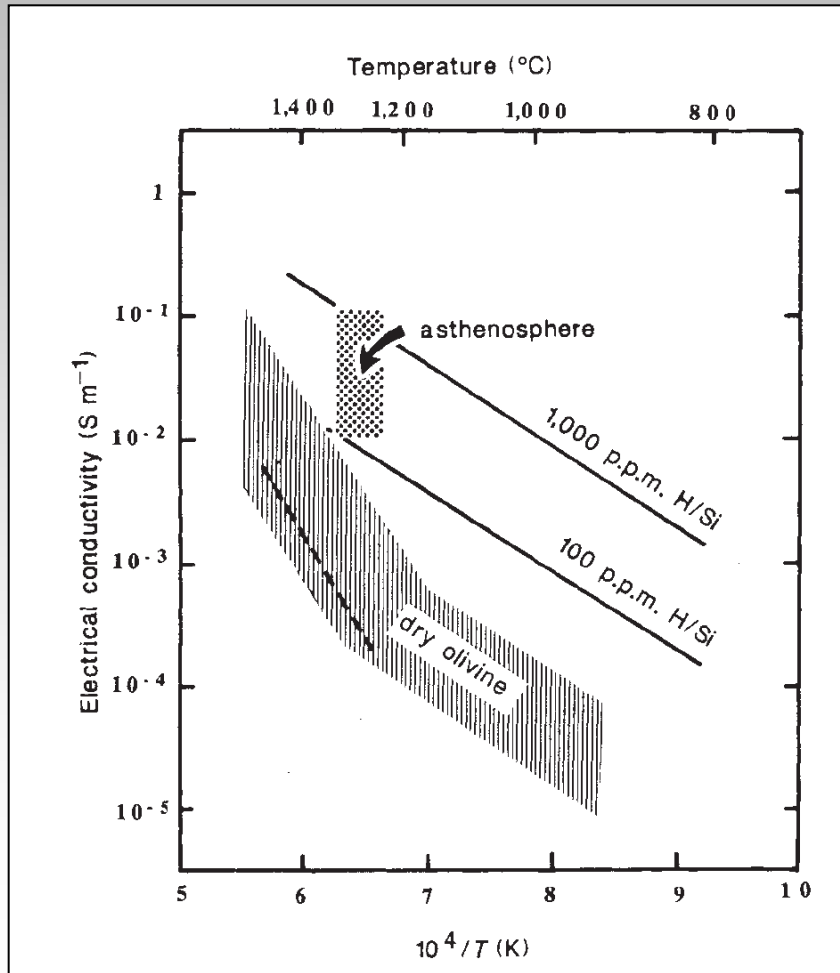
- (1) In regions with low geothermal gradient
- (2) Close to melting point

Figure 13.5

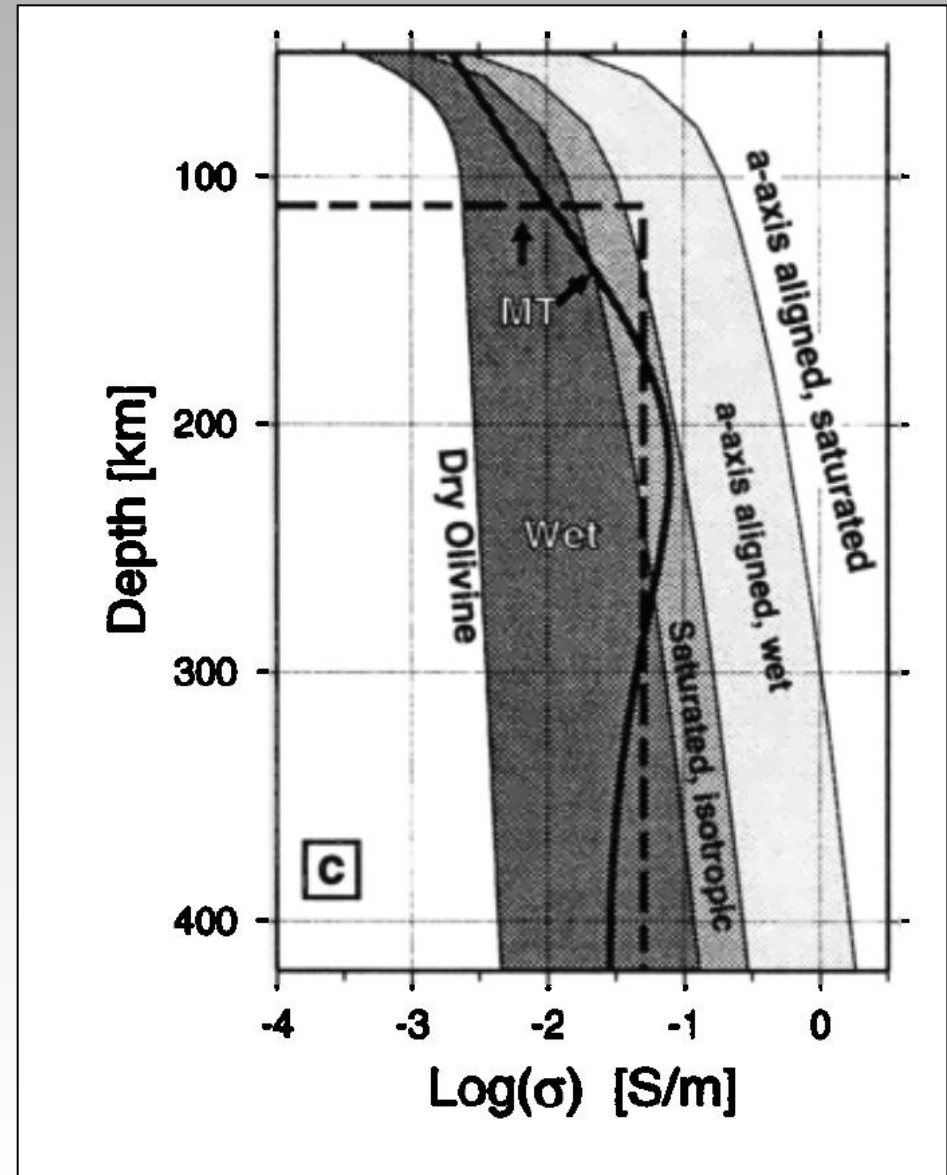
13.2 : Electromagnetic methods – water in nominally anhydrous minerals

The role of hydrogen in the electrical conductivity of the upper mantle

S. Karato

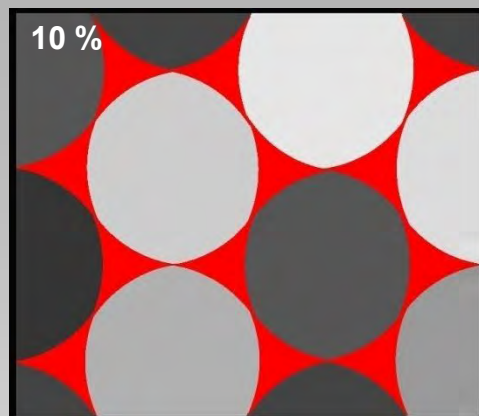


Karato, Nature, 1990

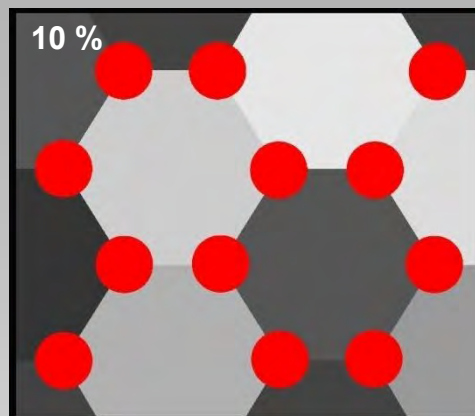


Lizzeralde et al., JGR, 1995, Figure 10c

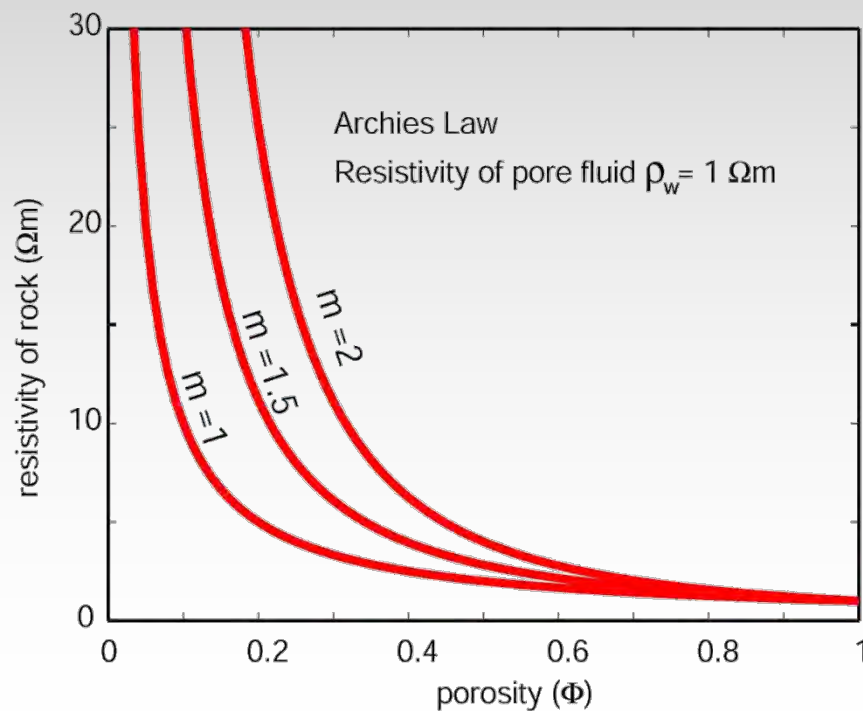
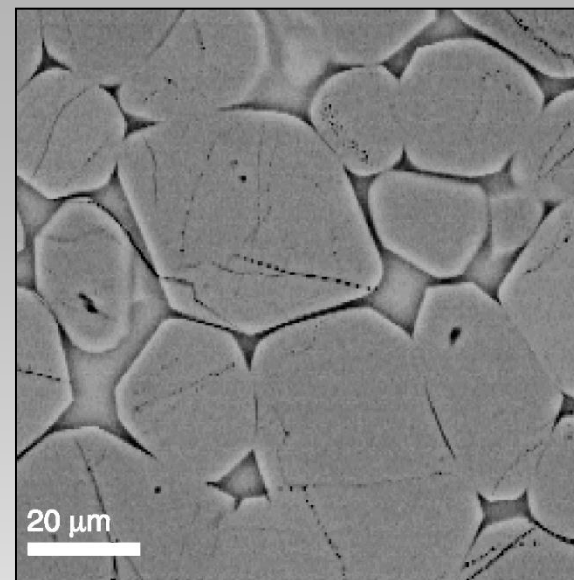
13.2 : Electromagnetic methods – partial melt



Dihedral angle = 0°



Dihedral angle = 180°

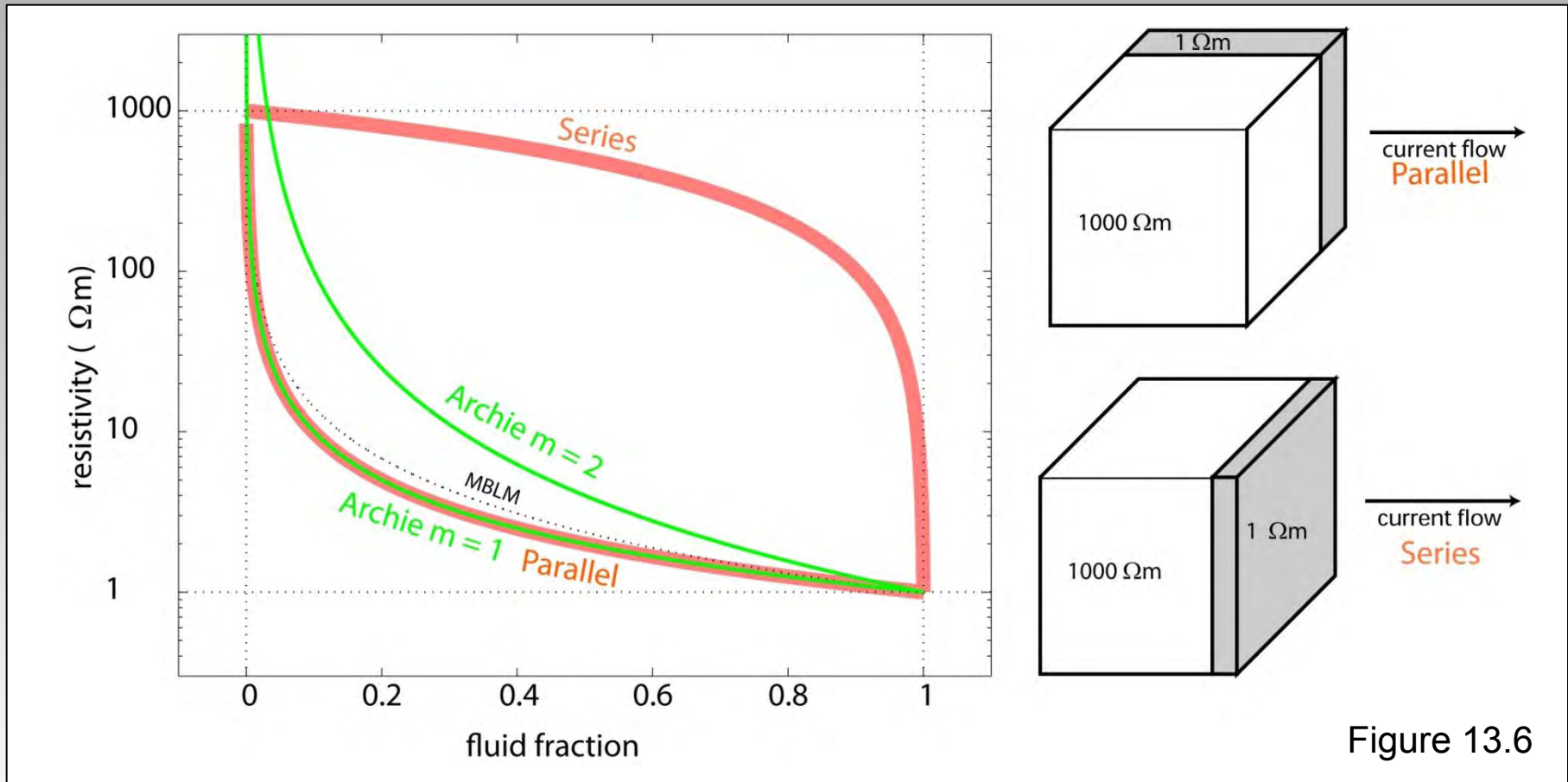


Olivine partial melt
ten Grotenhuis et al., JGR (2005)
10% melt

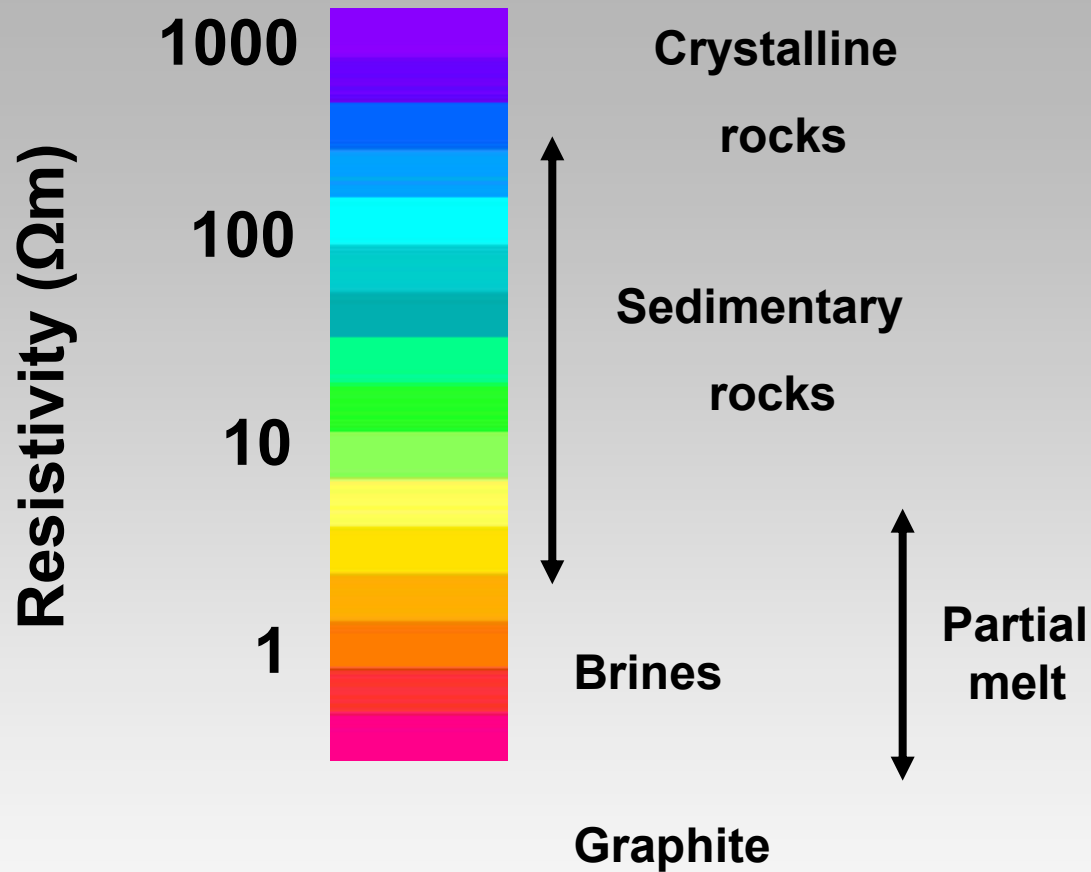
$m = 1.3$ in Archie's Law

Melt resistivity = $1 - 0.1 \Omega\text{m}$

13.2 : Electromagnetic methods – bulk resistivity of rocks containing fluids



13.2 : Electromagnetic methods – bulk resistivity of rocks containing fluids



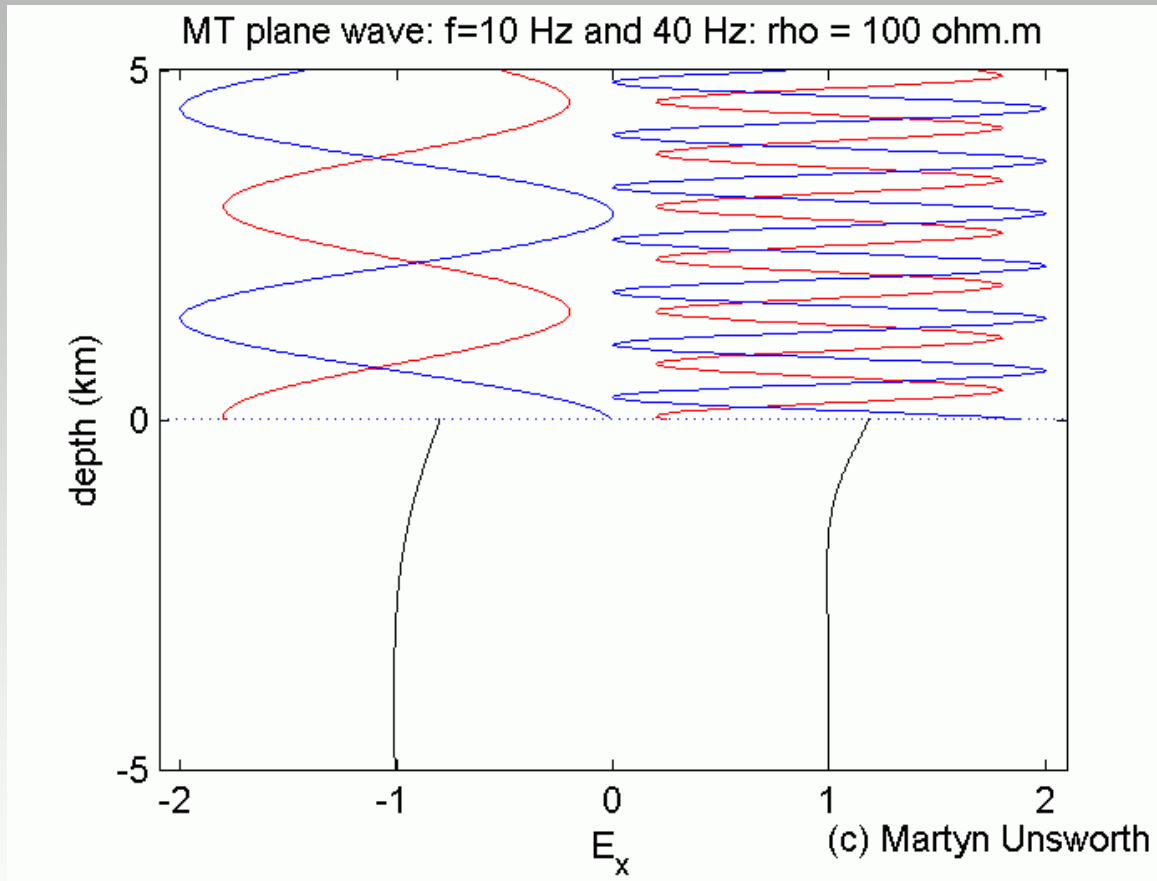
Beware of non-uniqueness in geophysics!

Low resistivity : brine, melt, graphite, sulphides

High conductivity = low resistivity

13.2 : Measuring the electrical resistivity of the Earth at depth (magnetotellurics)

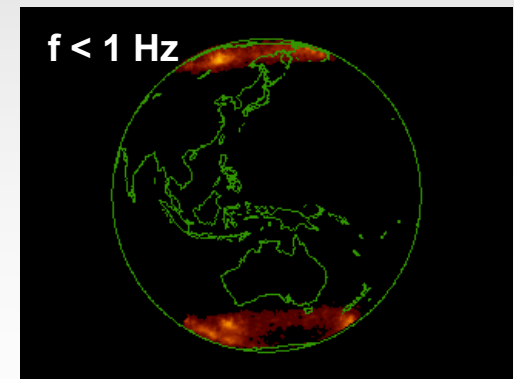
Figure 13.7



- Ratio of electric and magnetic fields $>$ resistivity
- Depth of investigation varies as $1/\sqrt{\text{frequency}}$



Global lightning activity



Aurora Borealis and Australis

13.2 : Measuring the electrical resistivity of the Earth at depth (magnetotellurics)

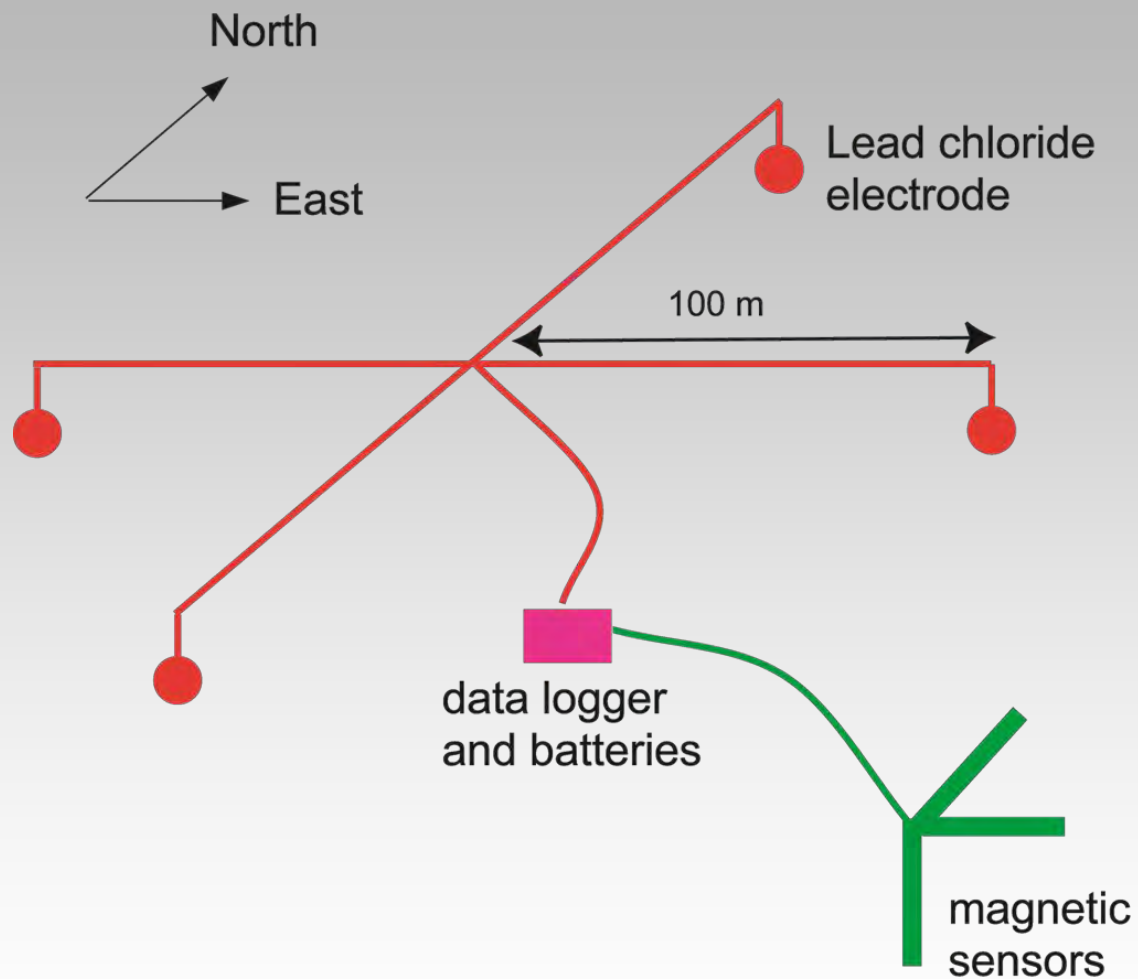


Figure 13.12

13.3 : Seismic methods - free water

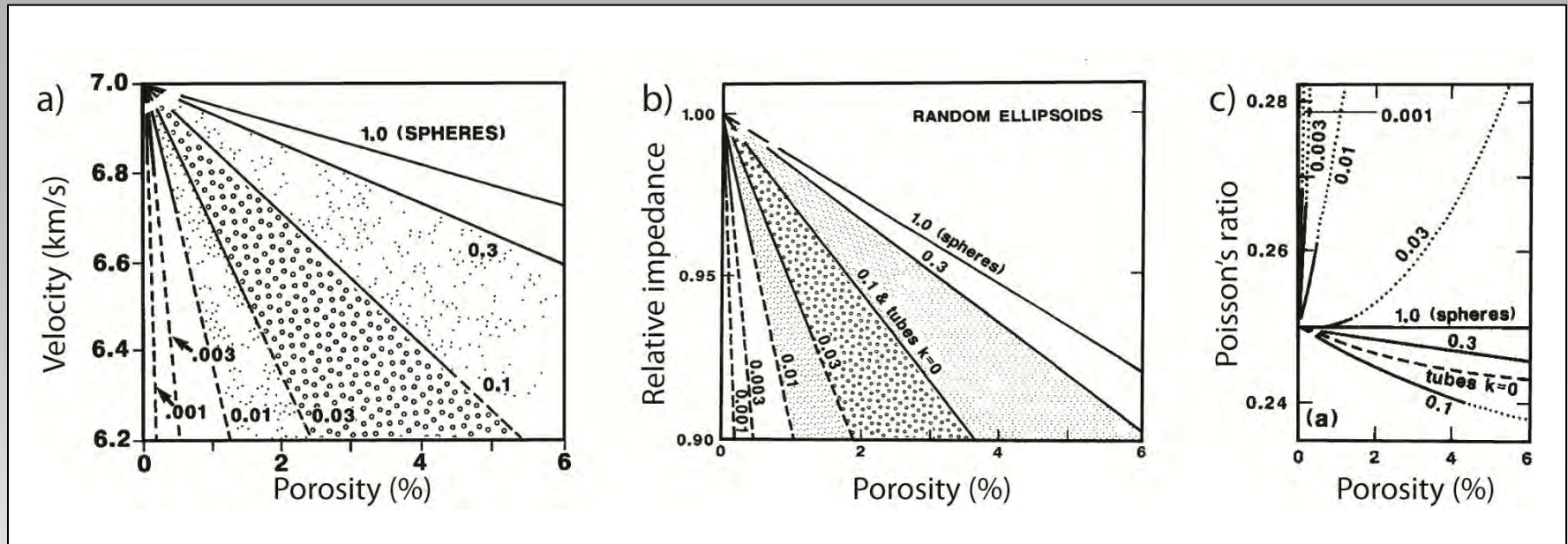


Figure 13.13
Hyndman and Shearer (1993)

Effect of free water is to

- Reduce the stiffness of the rock and lower velocity
- Increase or reduce Poisson's ratio (depending on aspect ratio of pores)
- Increase attenuation (grain boundary sliding, liquid squirt)
- Cause anisotropy

13.3 : Seismic methods - hydrous minerals

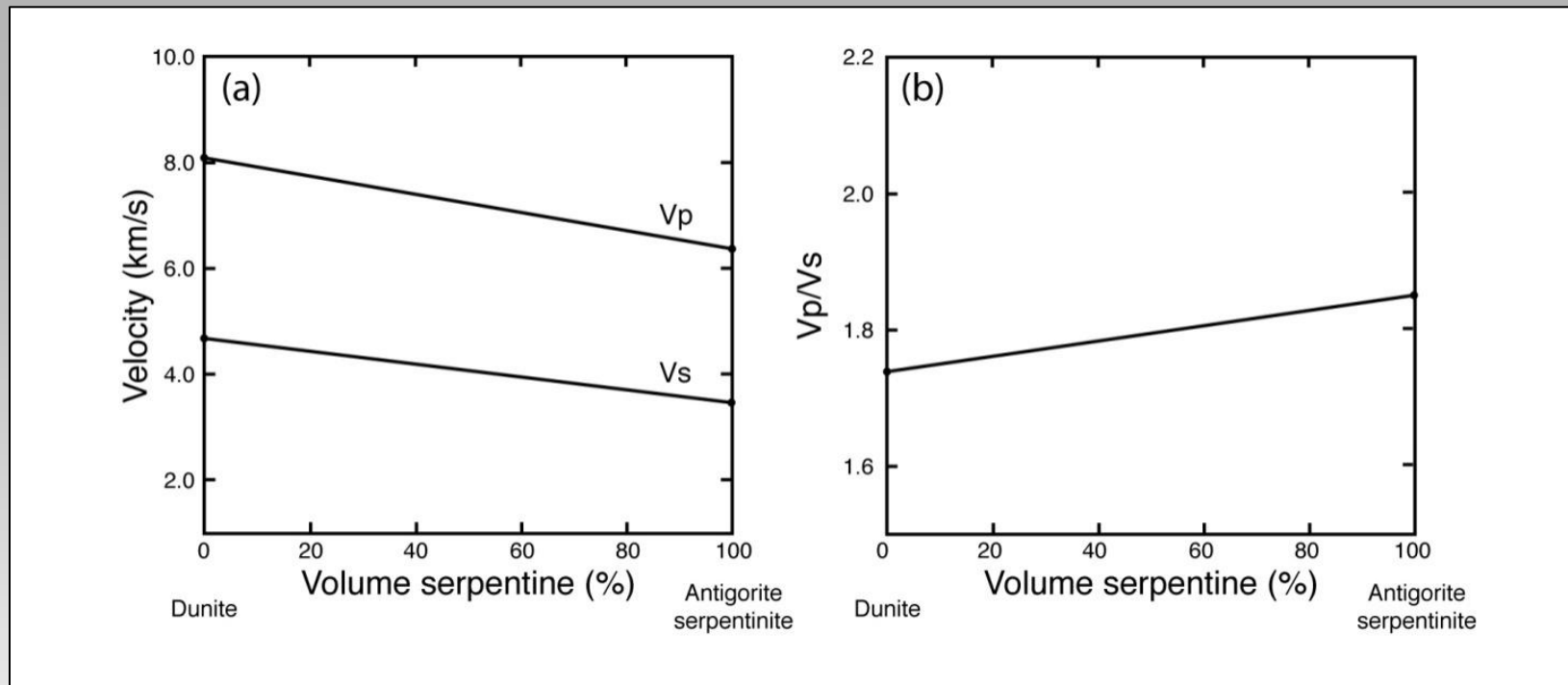


Figure 13.15

- contain water as part of chemical structure
- generally lower velocity than equivalent anhydrous mineral
- often anisotropic (antigorite P-wave = 71%; S-wave = 68%).
- macro anisotropy requires orientation by deformation

13.3 : Seismic methods - water in nominally anhydrous minerals

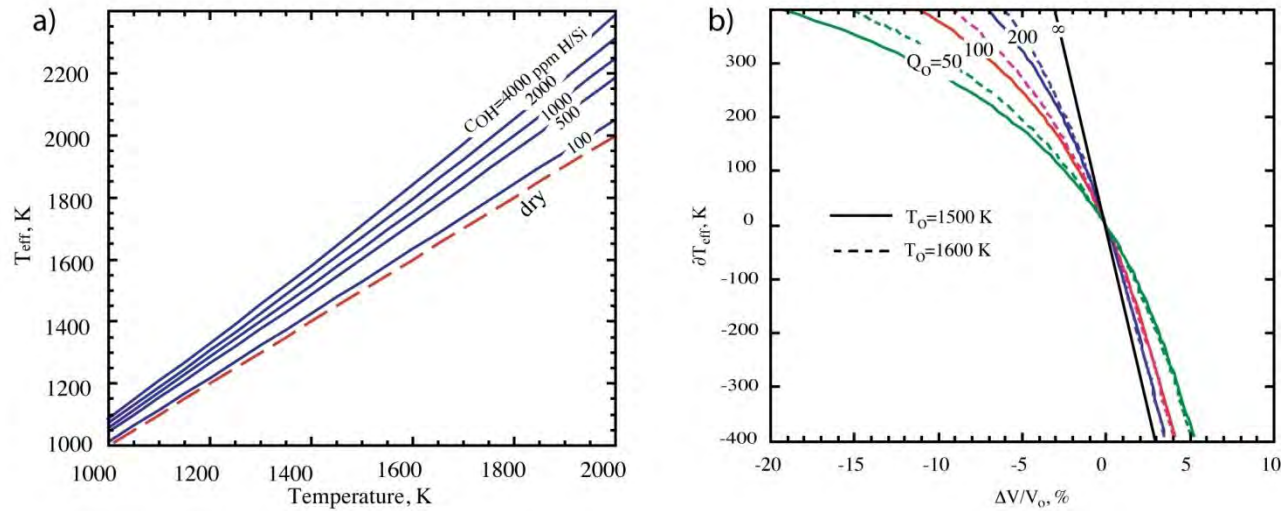
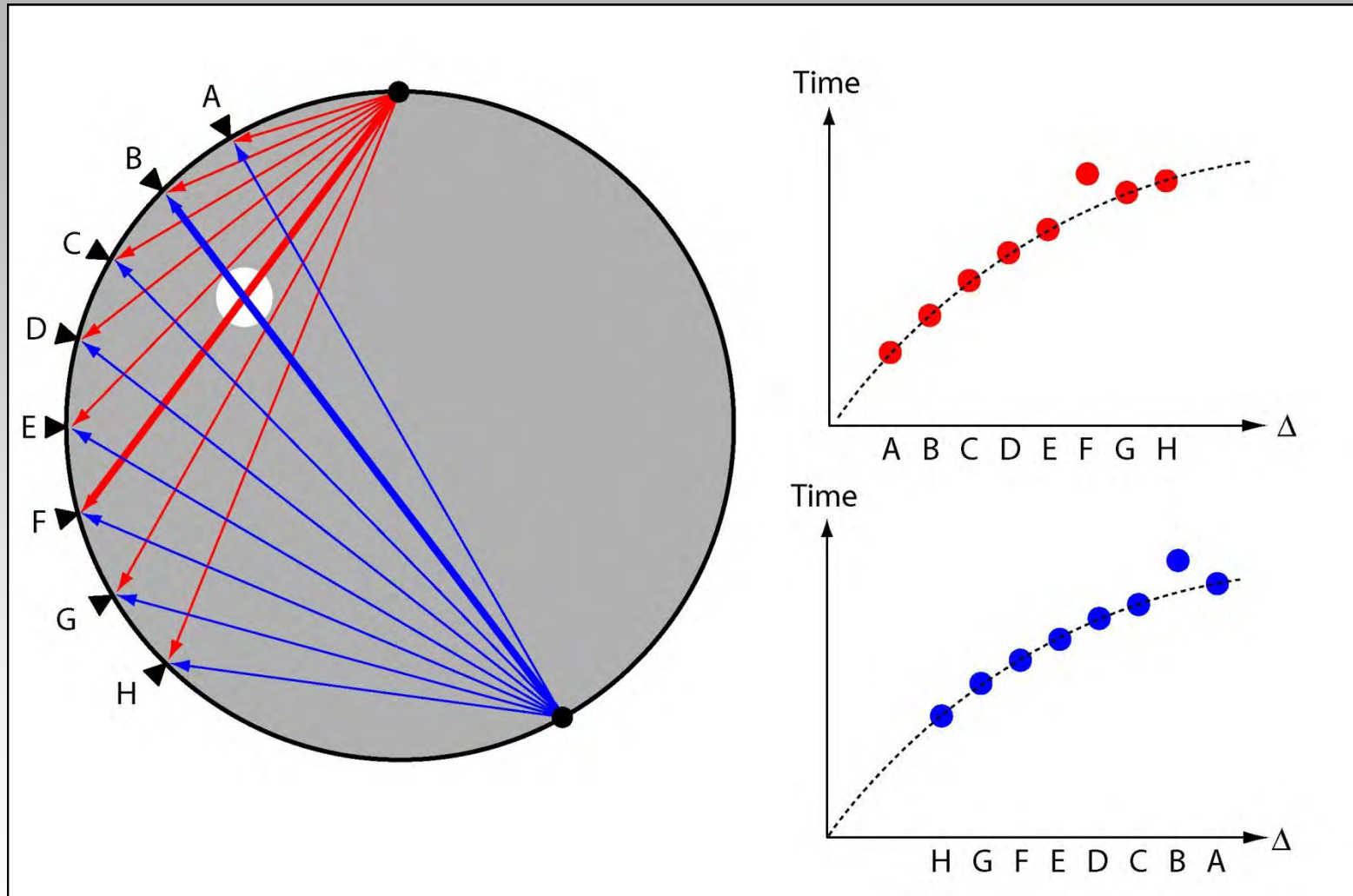


Figure 13.16

- H+ point defects in olivine
- Enhancement of anelasticity, modifies velocities
- Causes different types of olivine fabrics, development of seismic anisotropy,

13.3 : Seismic methods for imaging fluids in the deep crust and upper mantle

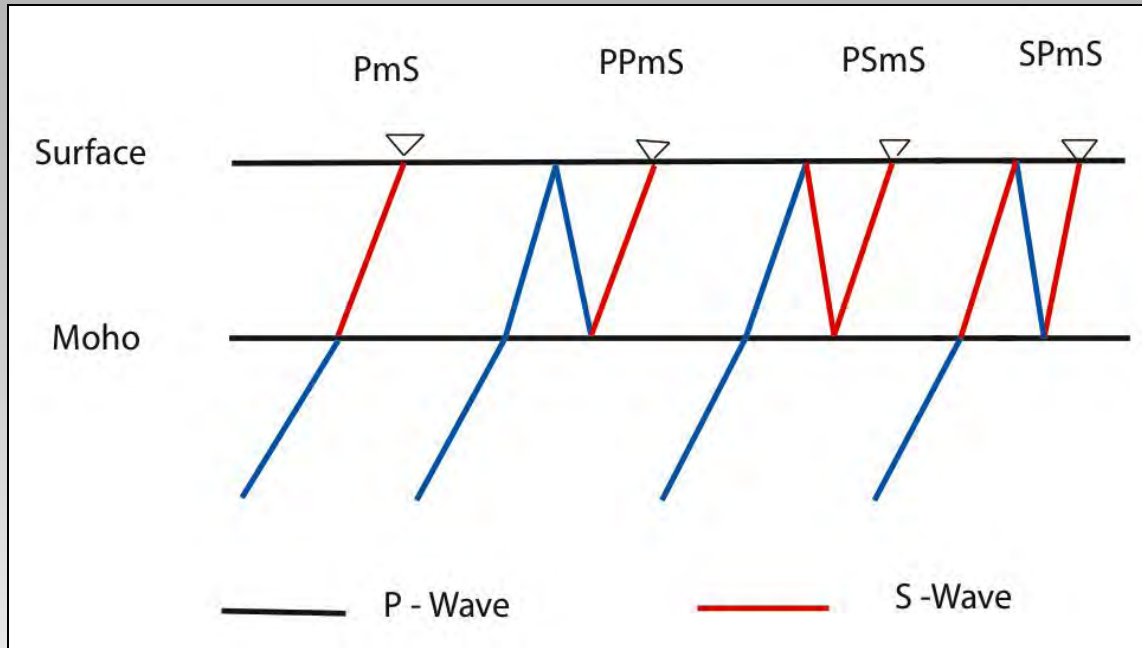
Seismic tomography



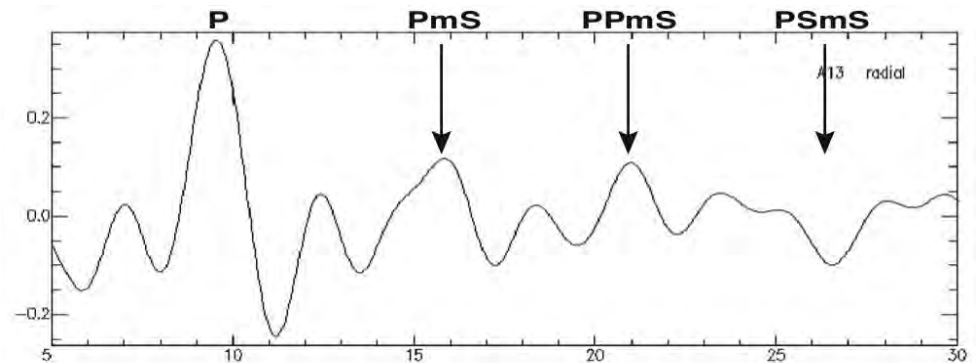
Energy sources : earthquakes, explosions
Can use travel times and waveform inversion
Fully 3-D approaches from large arrays

13.3 : Seismic methods for imaging fluids in the deep crust and upper mantle

Receiver function analysis



**Beware of
deceiver
functions !!!**



13.3 : Seismic methods for imaging fluids in the deep crust and upper mantle

Shear wave splitting – detect anisotropy

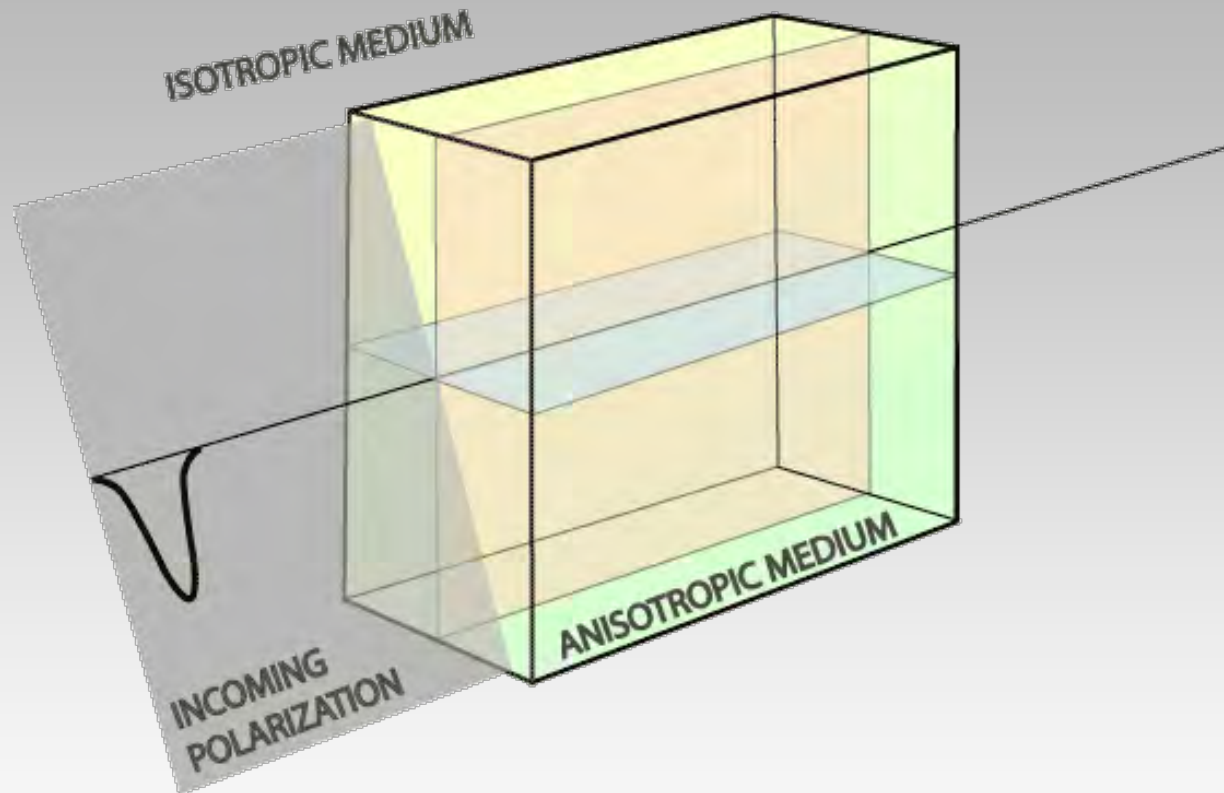
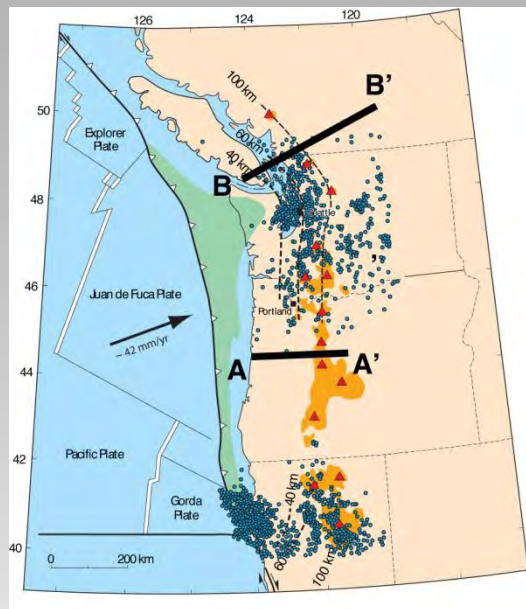
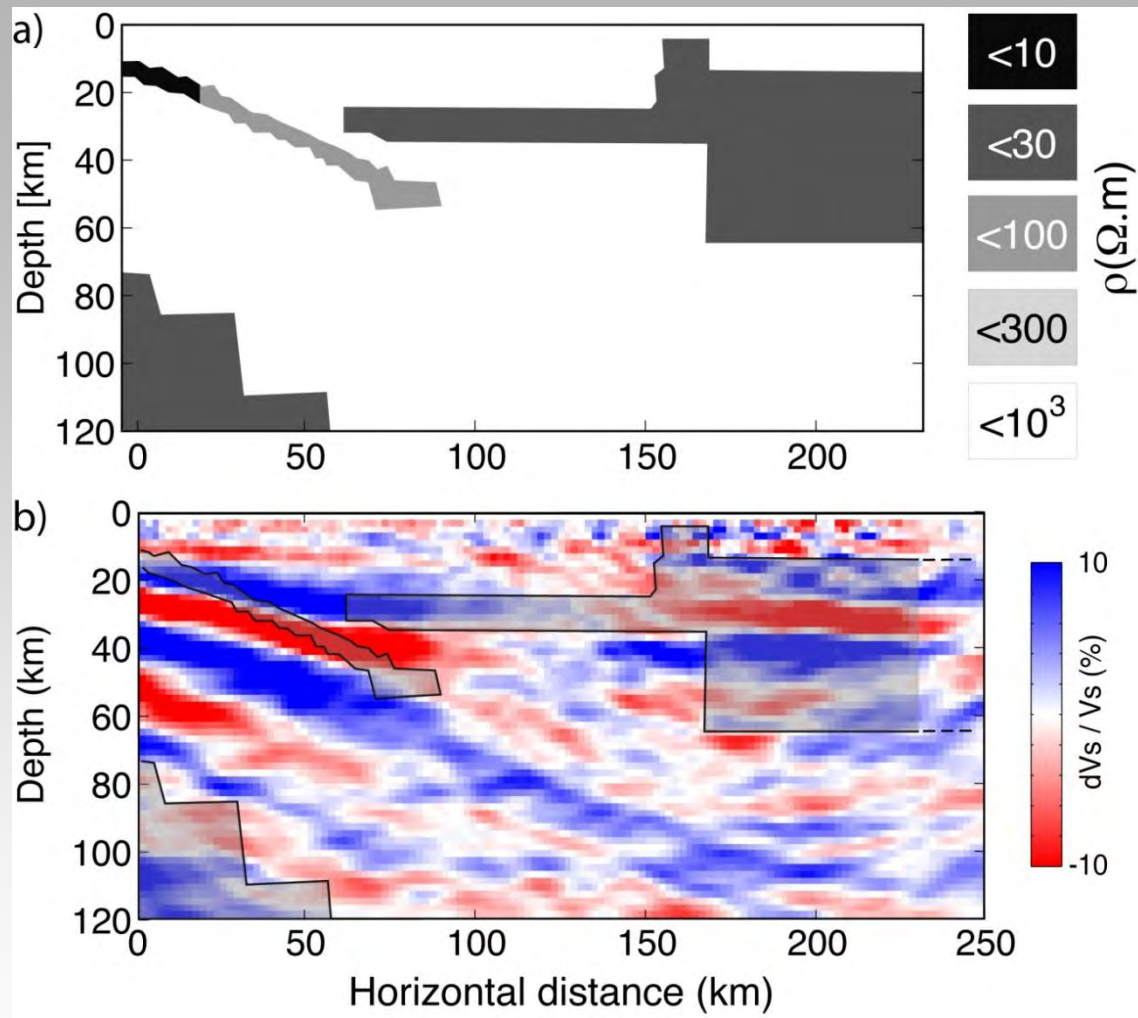


Figure by Ed Garnero, ASU

13.5 : Subduction zones - Cascadia



JDF plate is hot and young (6-10 Ma)

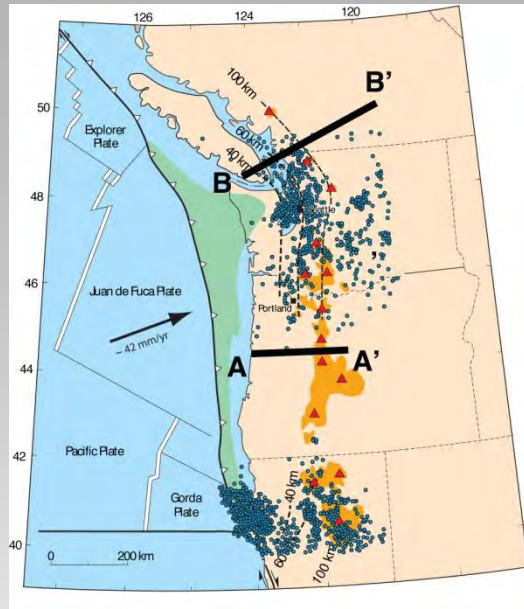


Oregon : Profile AA'

Magnetotellurics : Wannamaker et al., (1989)

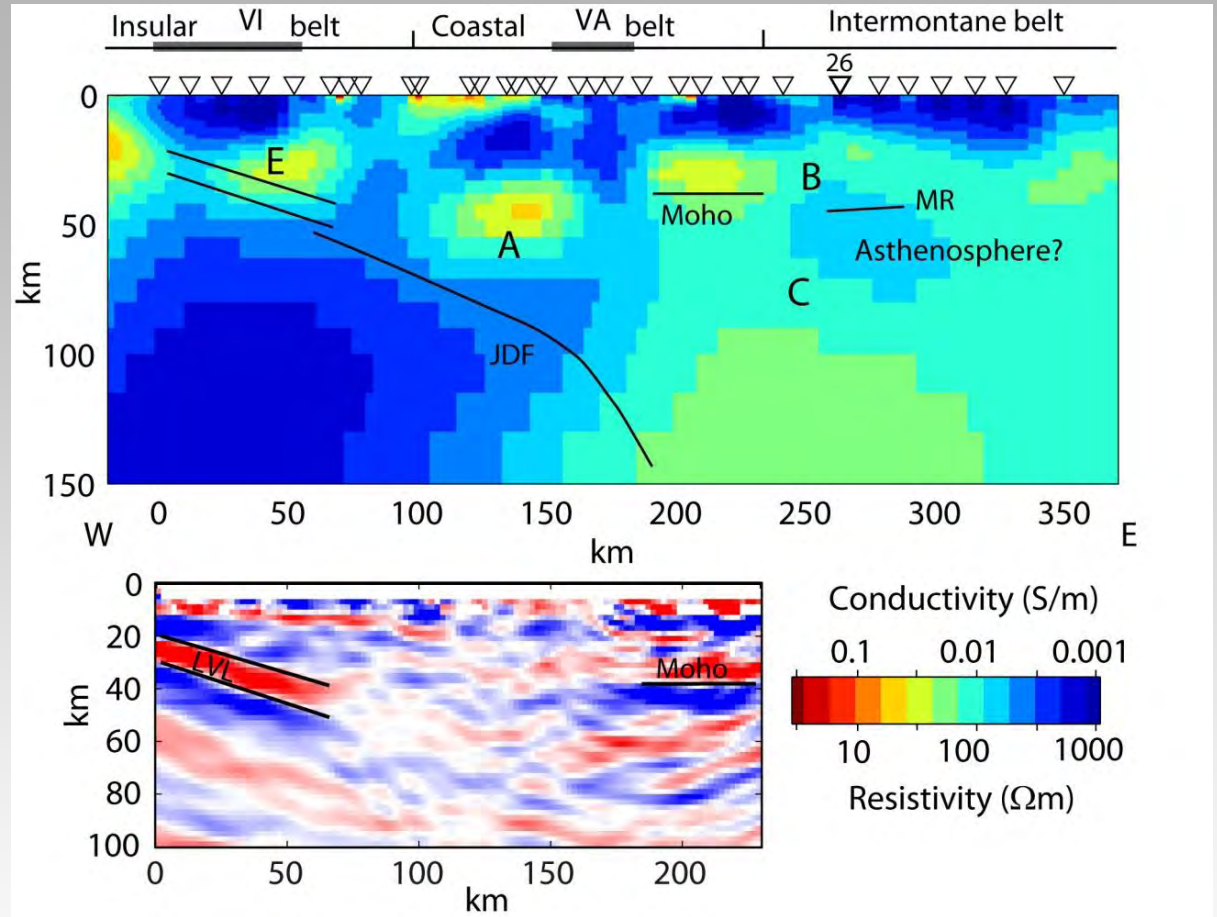
Seismic RF : Rondenay et al., (2001)

13.5 : Subduction zones - Cascadia



E : fluids expelled from plate into overlying crust
Channel in crust or along boundary?

A : Mantle wedge conductor. Antigorite lowers seismic velocity, but not enough. Free fluids?

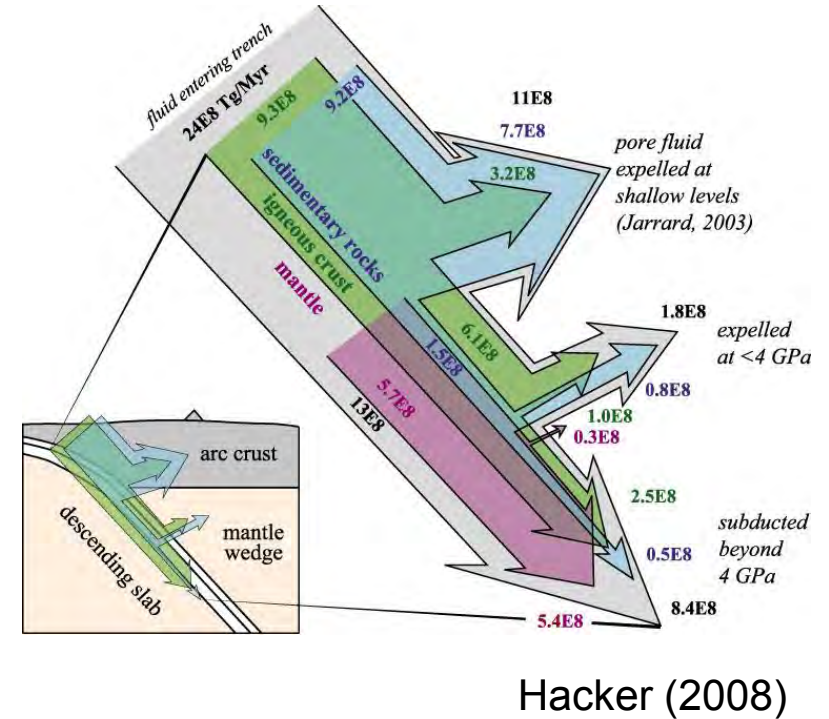
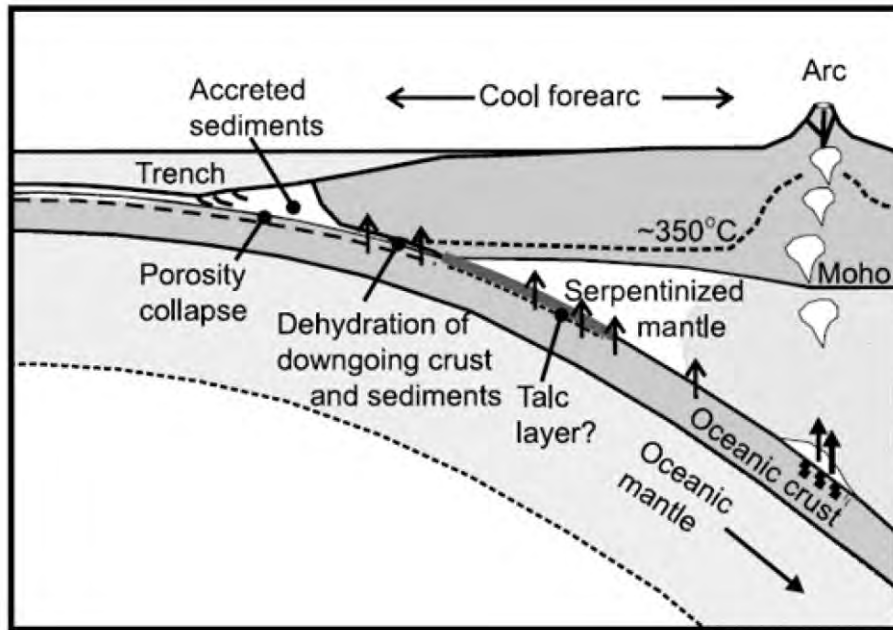


Southern British Columbia (BB')

Magnetotellurics : Soyer and Unsworth (2006)

Seismic RF : Nicholson et al., (2005)

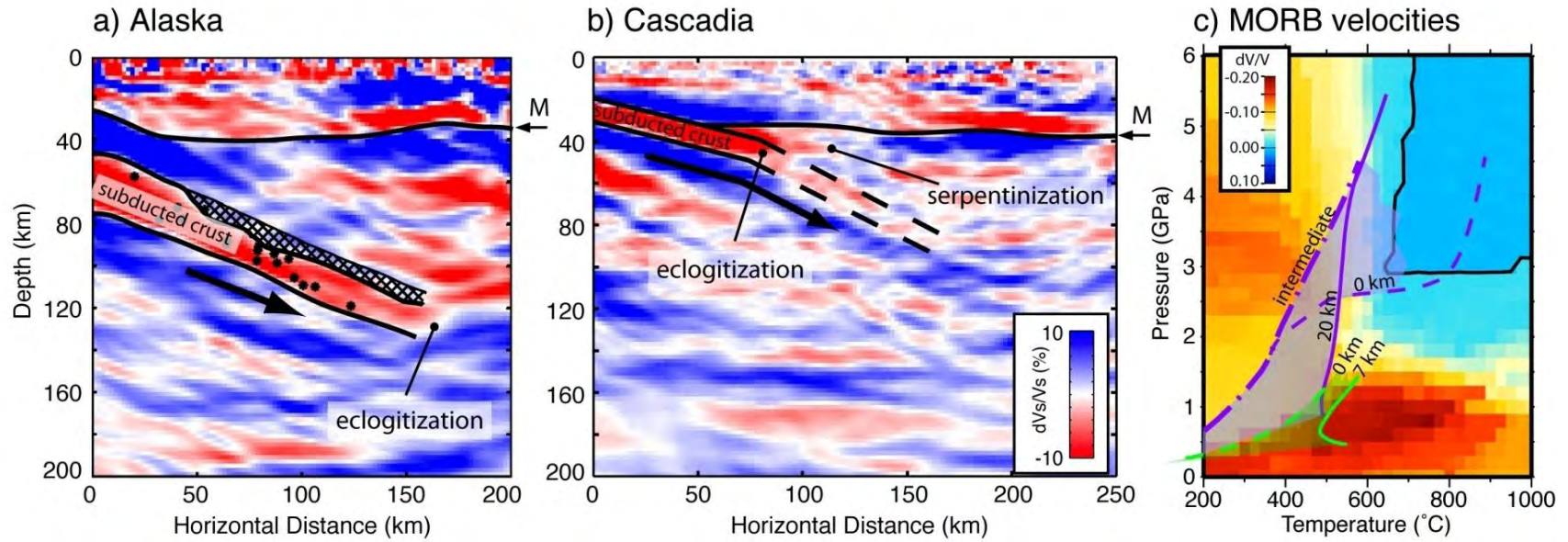
13.5 : Subduction zones



Water transported to depth as

- Free water in sediments and oceanic crust
- Hydrated minerals in oceanic crust and upper mantle (serpentine)

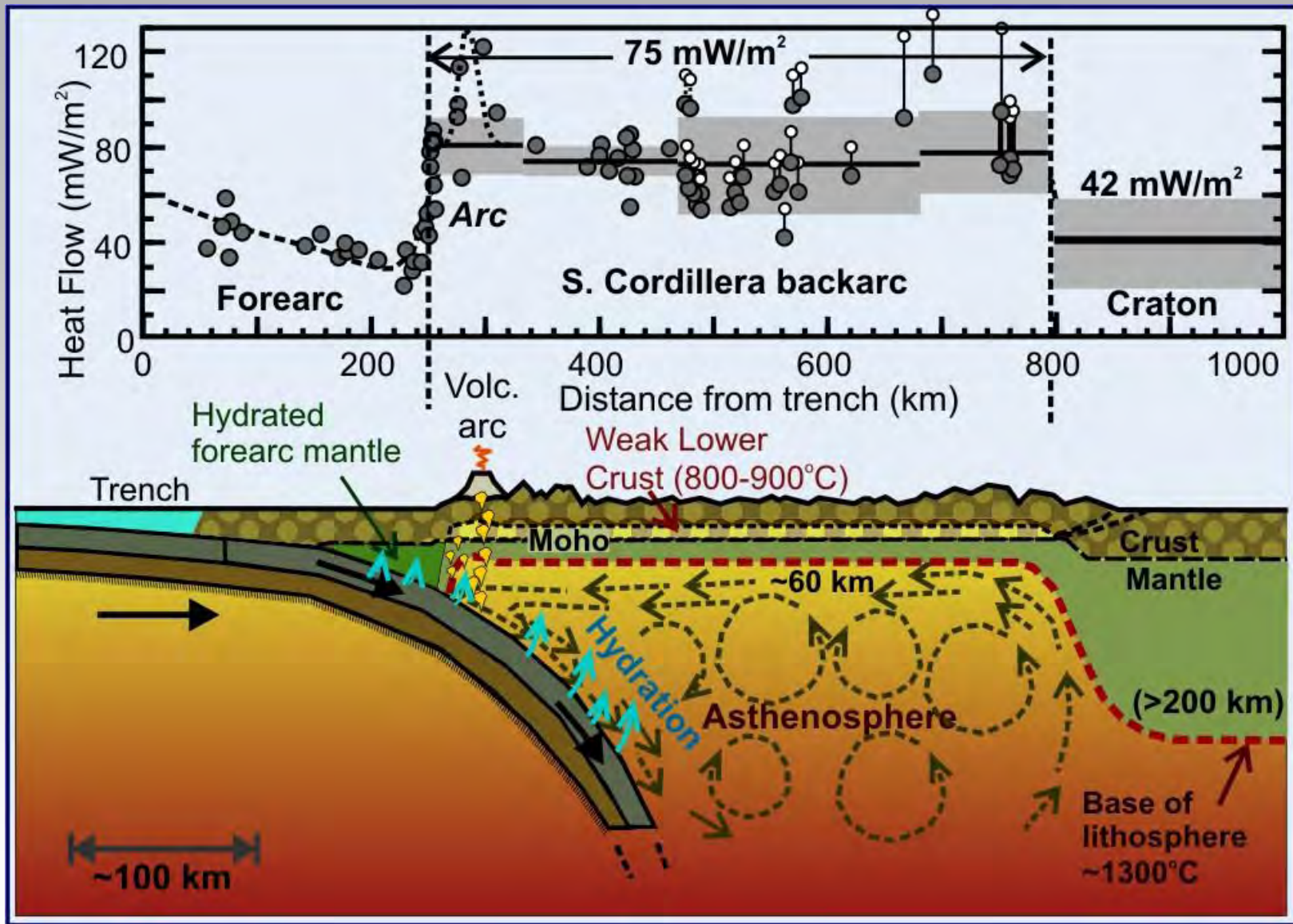
13.5 : Subduction zones



PT paths : **Alaska** **Cascadia**

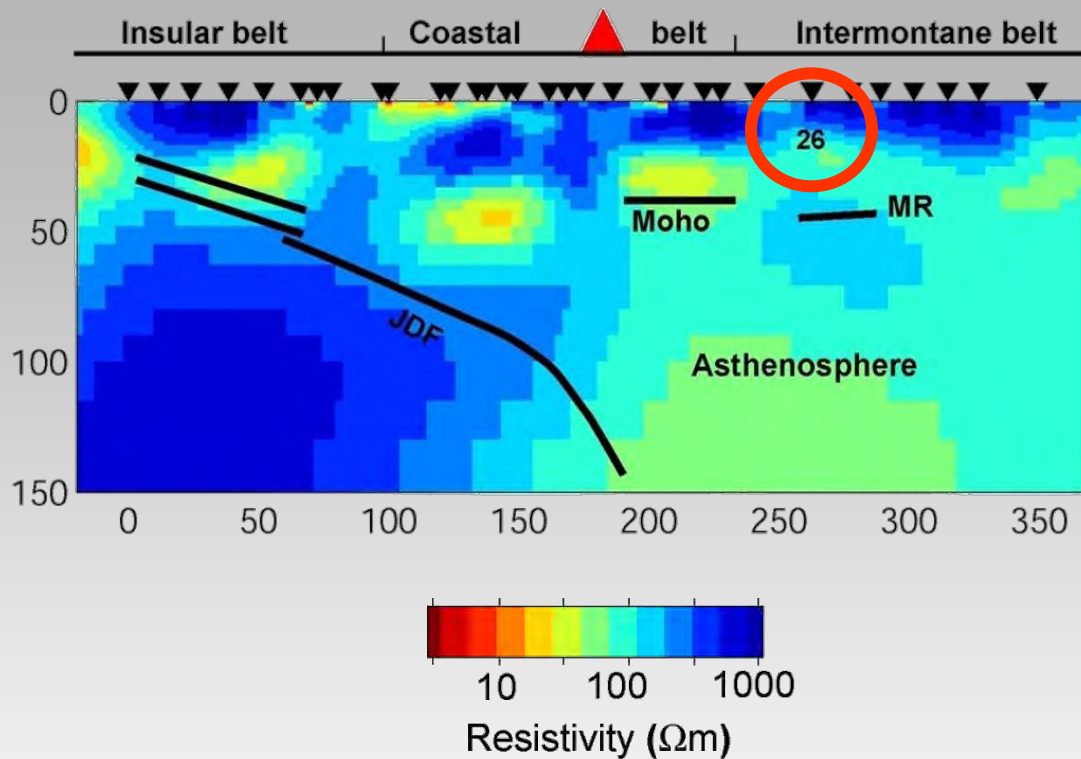
Rondenay et al., 2008

13.5 : Subduction zones - Cascadia

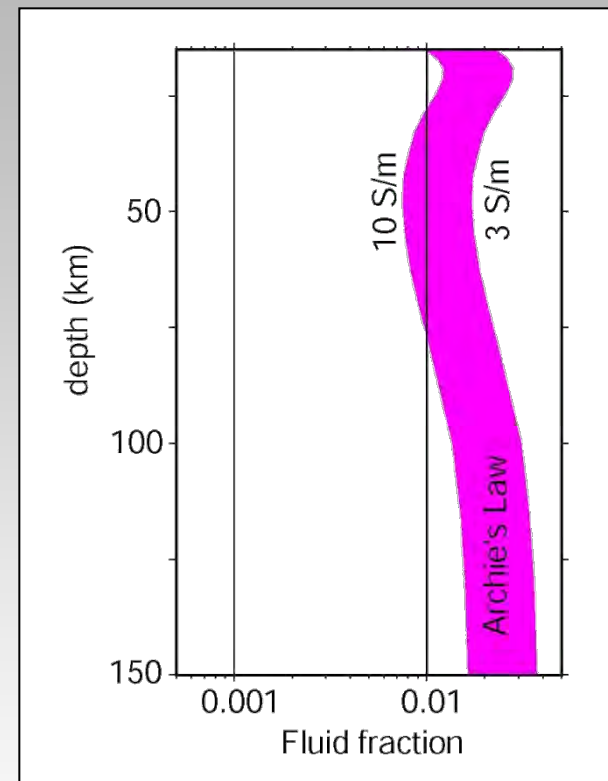


Hyndman et al., *GSA Today*, (2005)

13.5 : Subduction zones - Cascadia

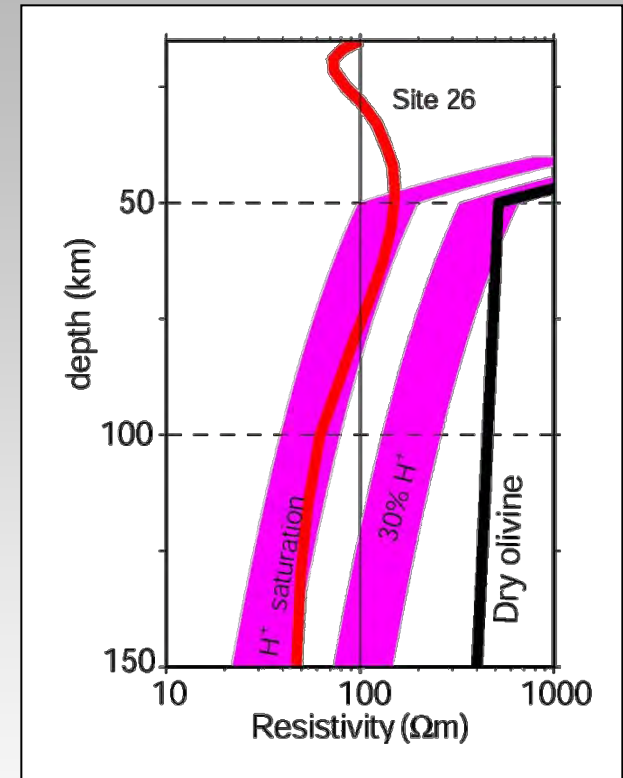
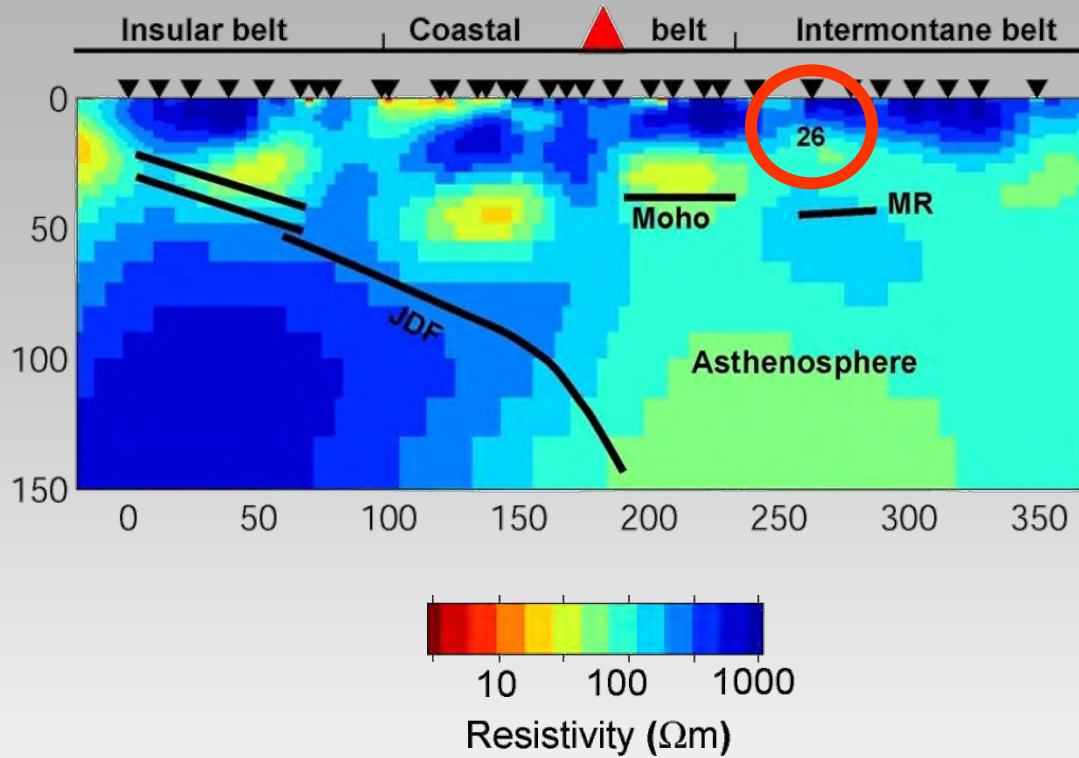


Soyer and Unsworth, *Geology*, 2006



Is the low resistivity in back arc due to melt / aqueous fluids?

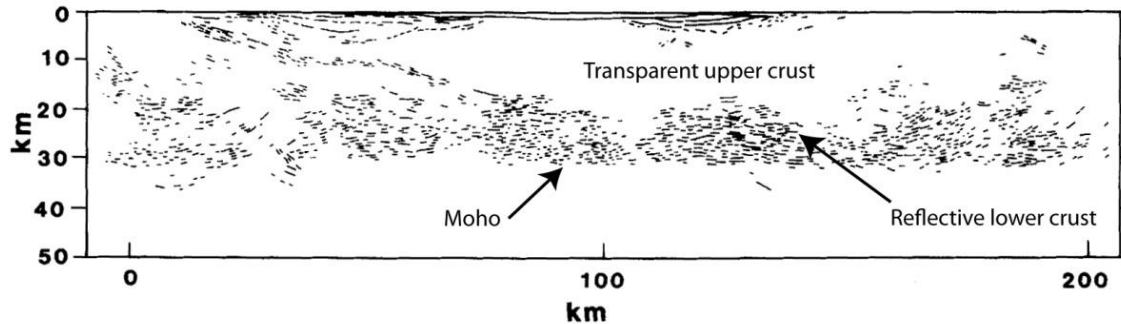
13.5 : Subduction zones - Cascadia



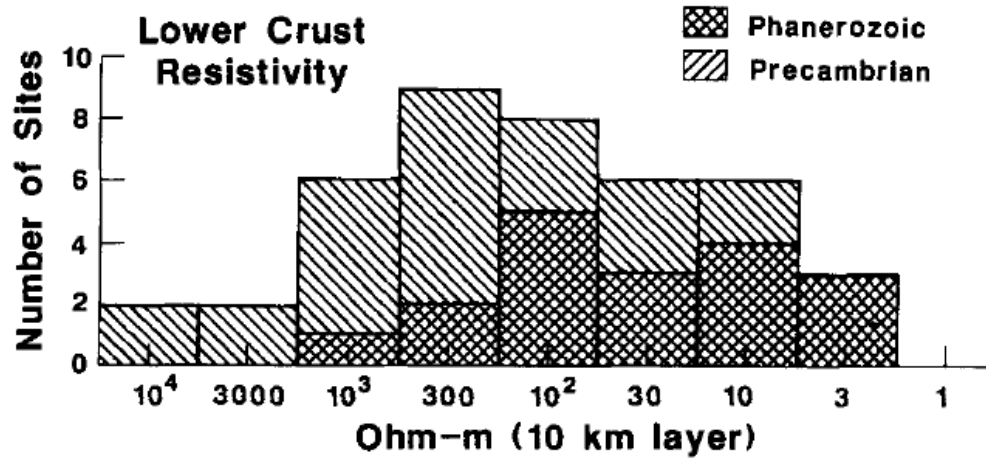
Soyer and Unsworth, *Geology*, 2006

Is the low resistivity in back arc due to hydrogen diffusion?

13.6 : Metasomatism in the stable continental lithosphere (Phanerozoic Crust)



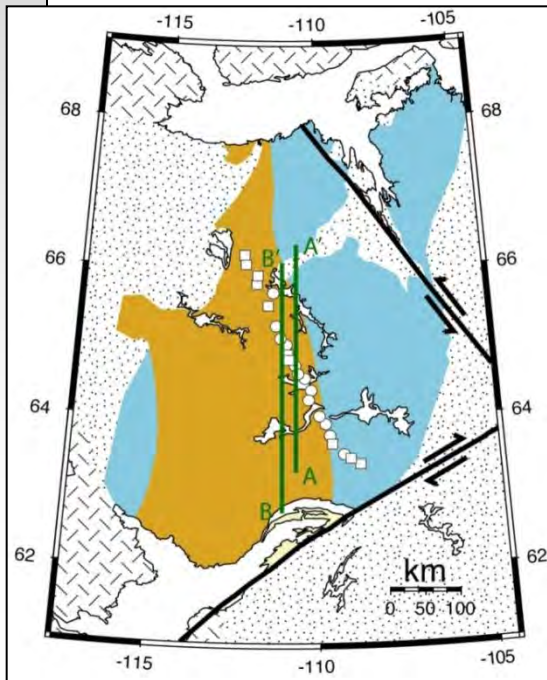
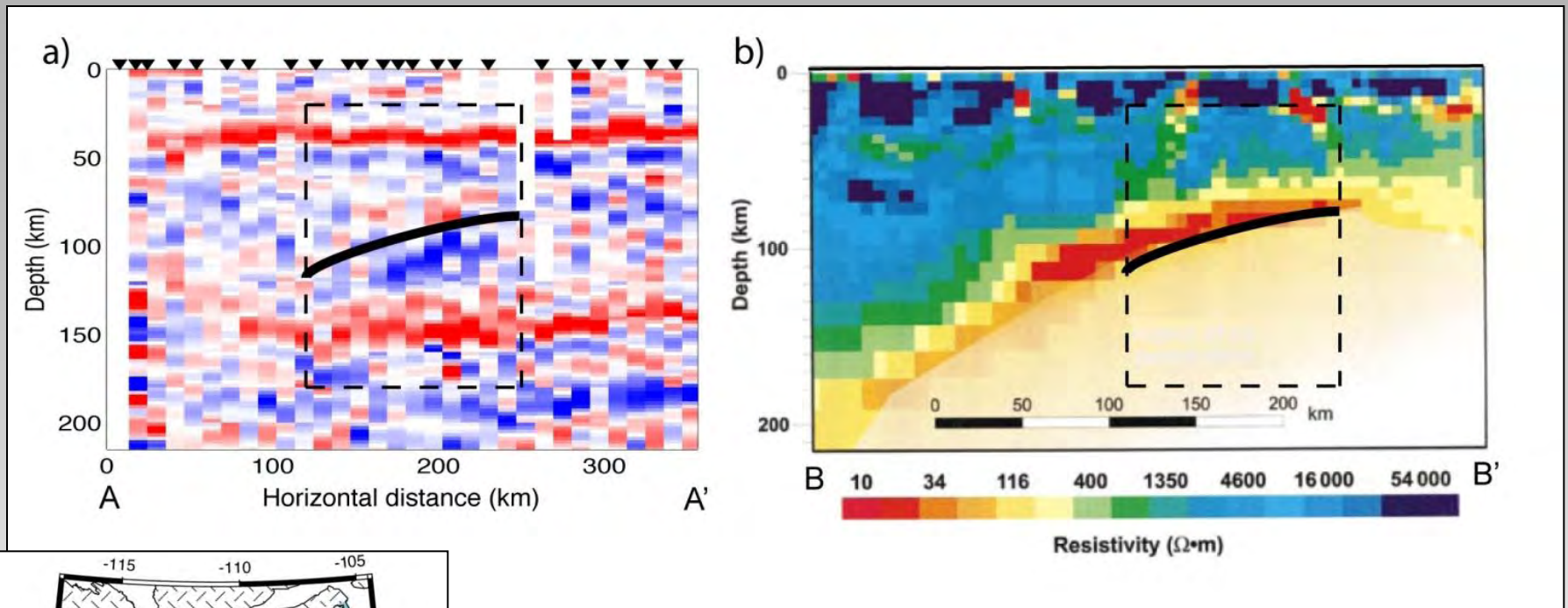
Matthews(1986)
Seismic reflections from
lower continental crust



Hyndman and Shearer (1989)

- Phanerozoic crust often characterized by
- Elevated conductivity in lower crust
 - Enhanced seismic reflectivity
 - Lower velocities than predicted by measurements on dry xenoliths

13.6 : Metasomatism in the stable continental lithosphere (Archean Cratons)



Slave Craton, Canada

Chen et al., (2009)

9-21% seismic velocity reduction – hydrated minerals

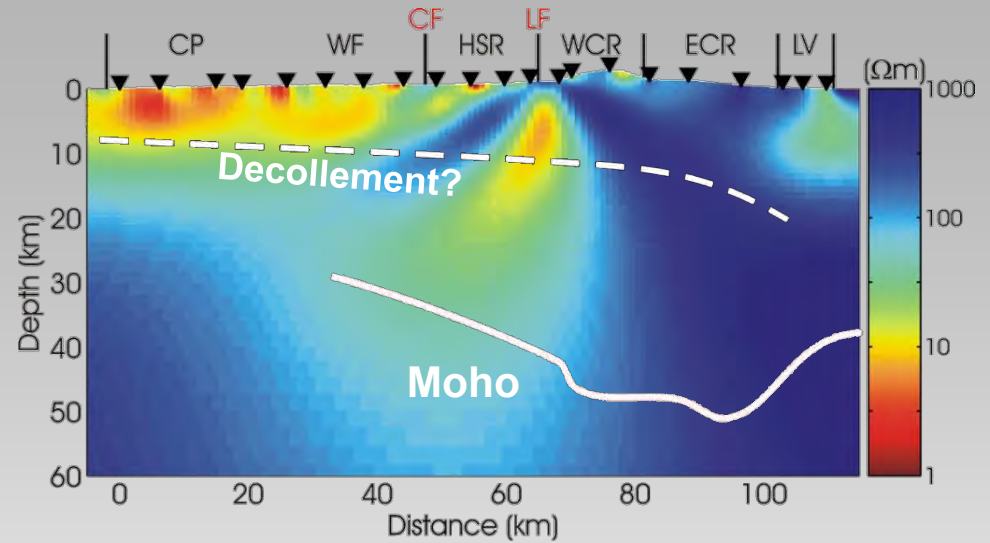
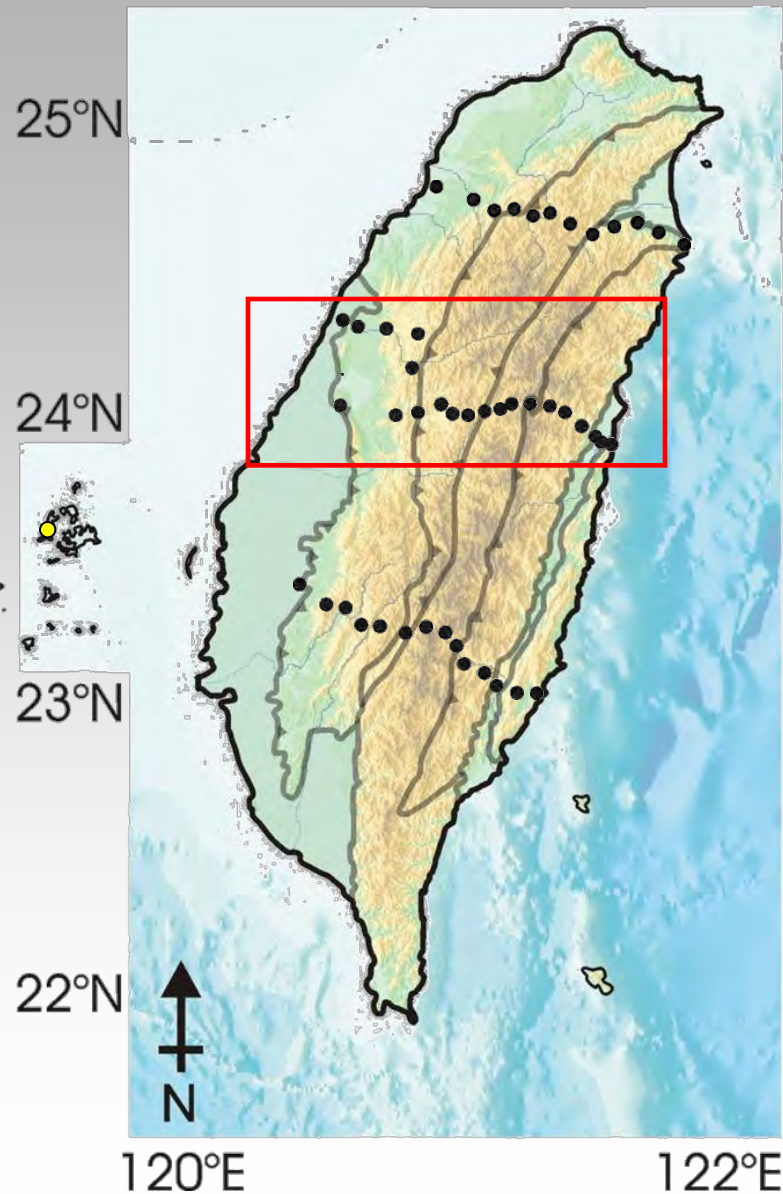
Resistivity decrease – graphite

Subduction event at 3.5 Ga

Jones and Ferguson (2001)

Decrease in resistivity at Moho (50000 – 5000 Ωm)

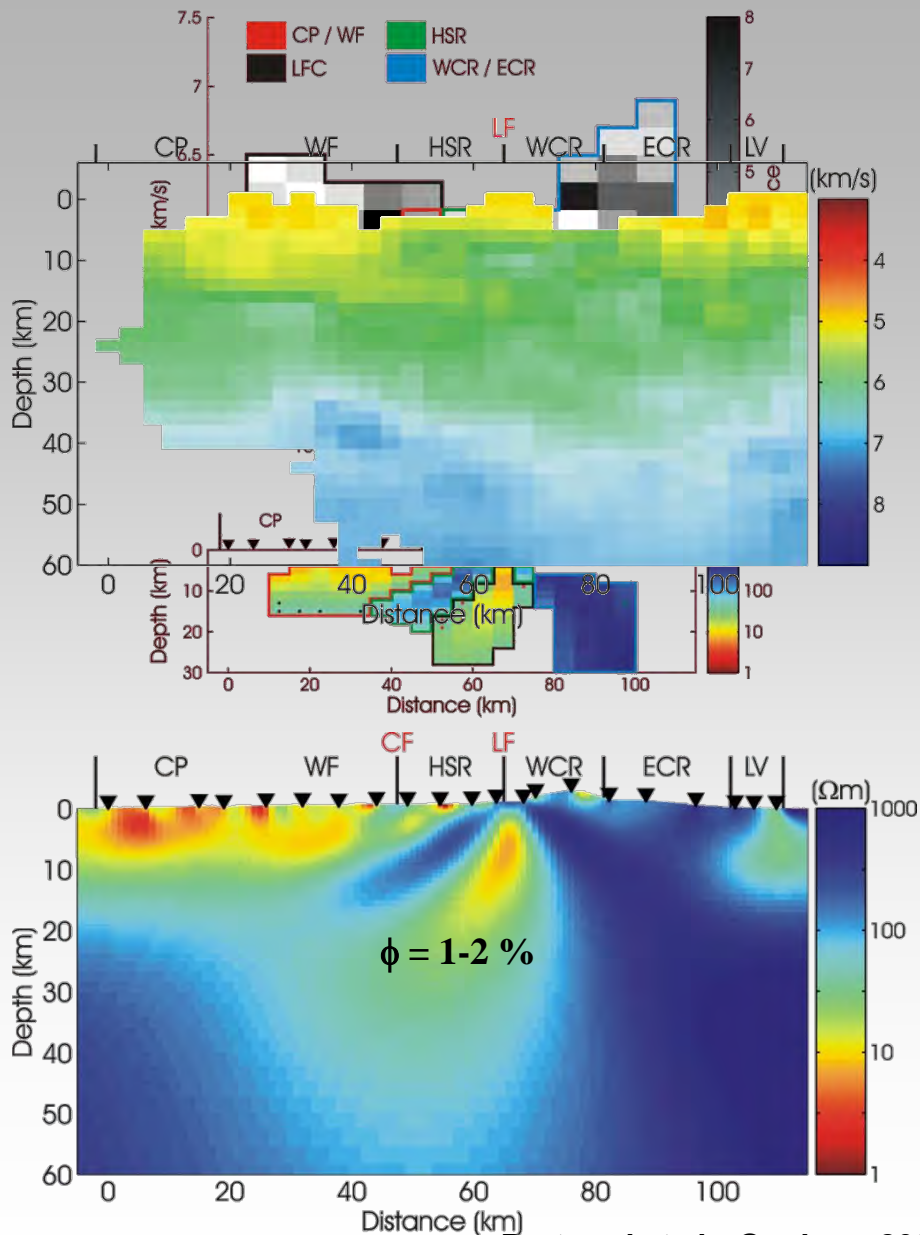
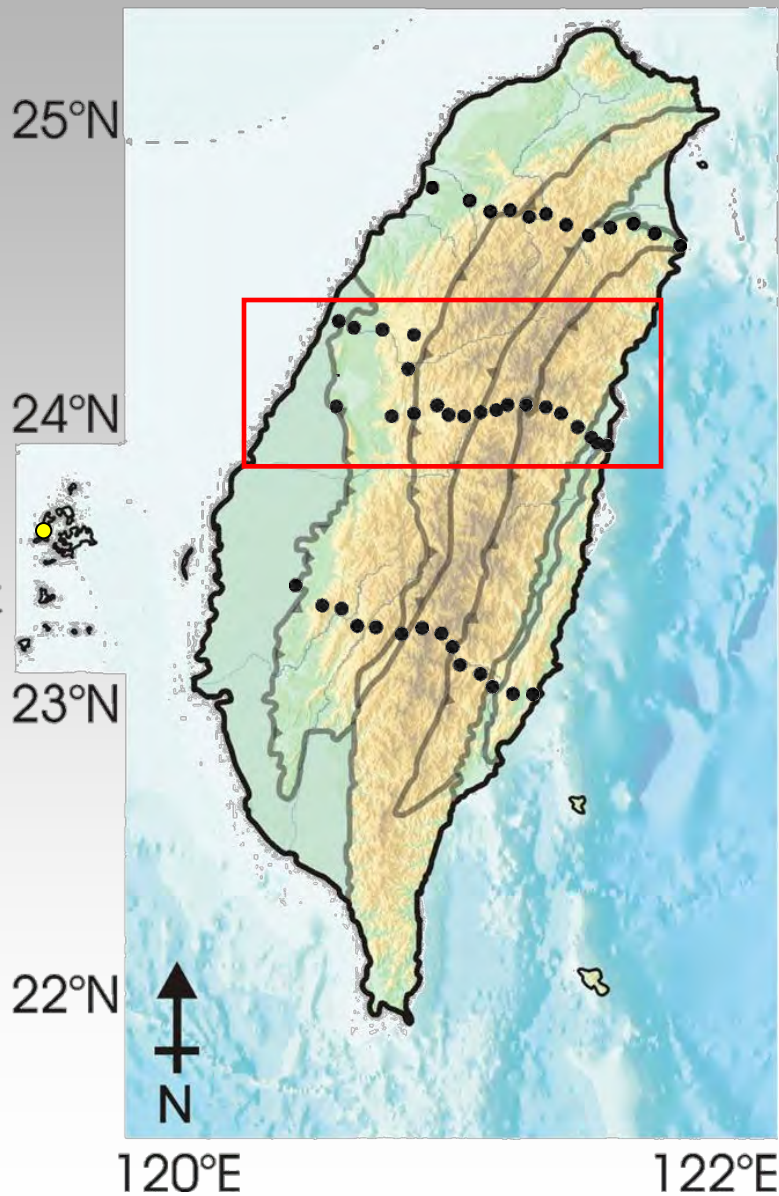
13.7 : Fluids generated in collision zones (arc-continent collision)



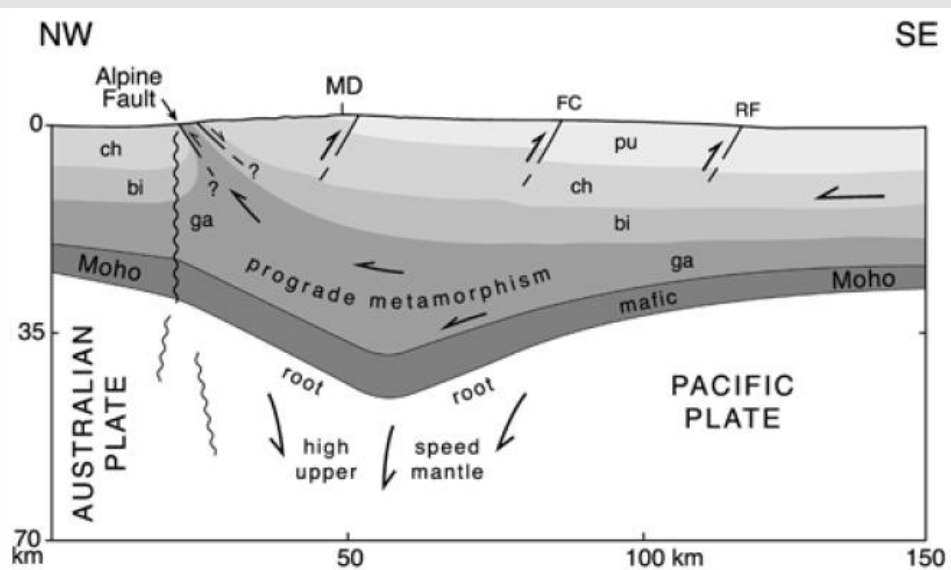
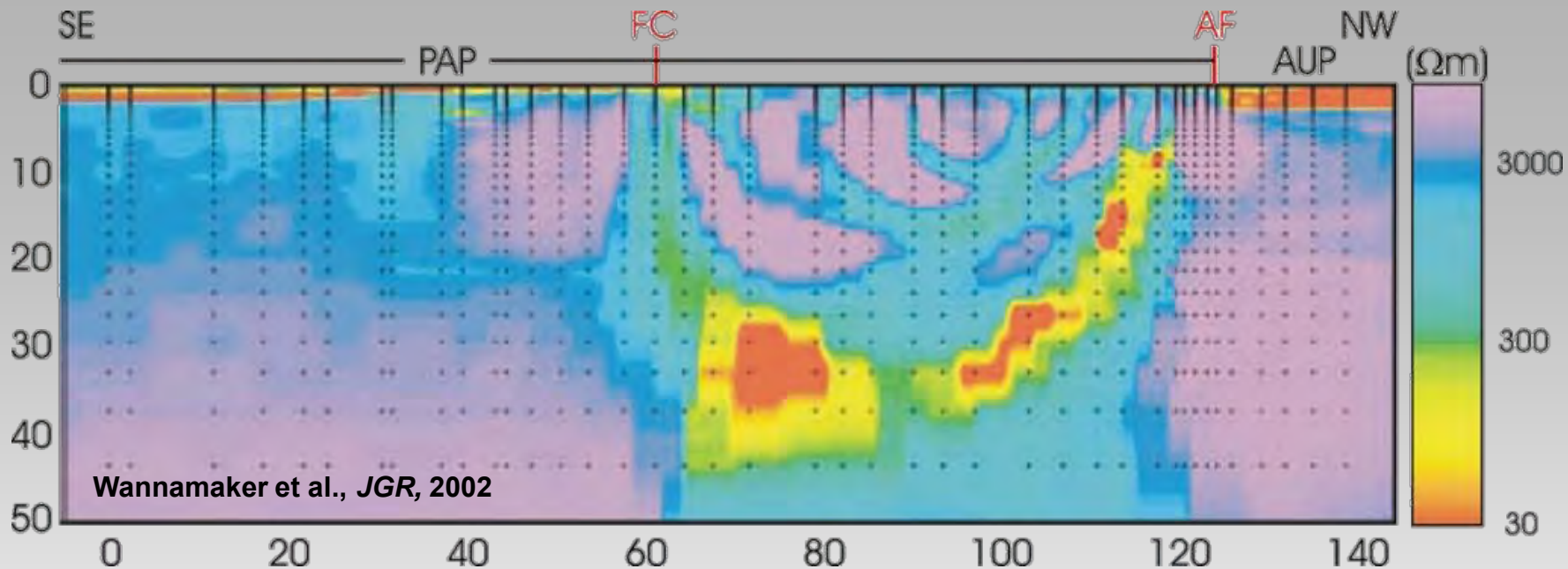
Bertrand et al., *Geology*, 2009

- Lishan fault conductor extends across inferred decollement
- Inconsistent with thin-skinned model
- Fluid originates in crustal root (prograde metamorphism)

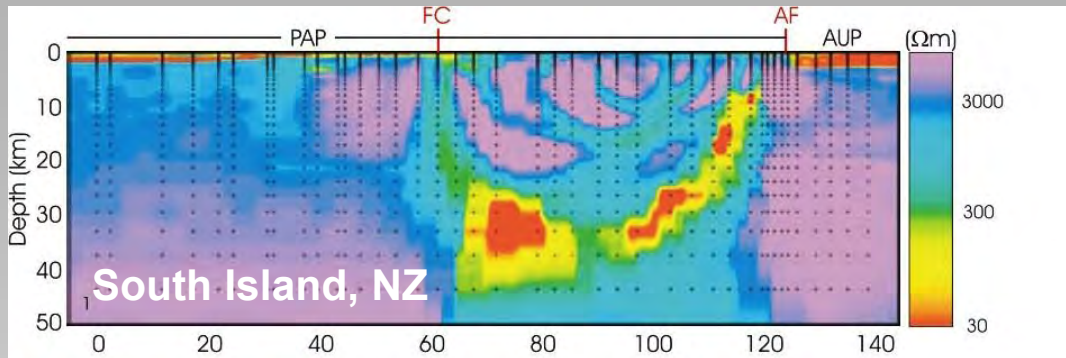
13.7 : Fluids generated in collision zones (arc-continent collision)



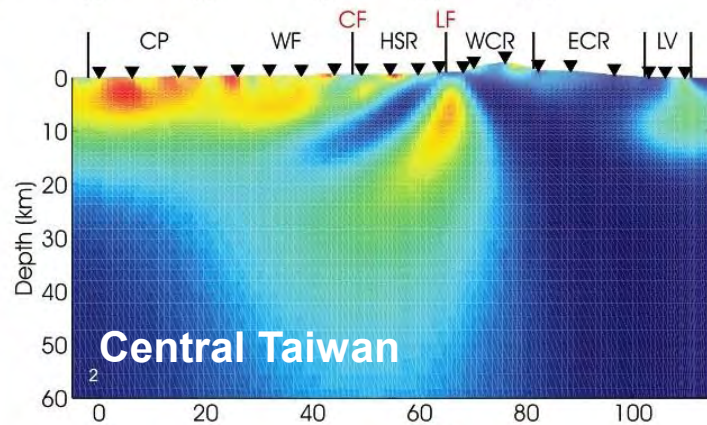
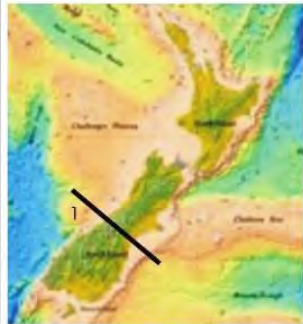
13.7 : Fluids generated in collision zones



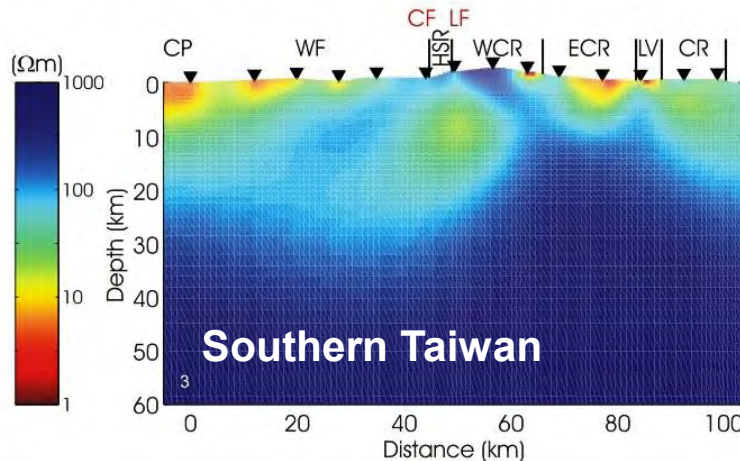
13.7 : Fluids generated in collision zones



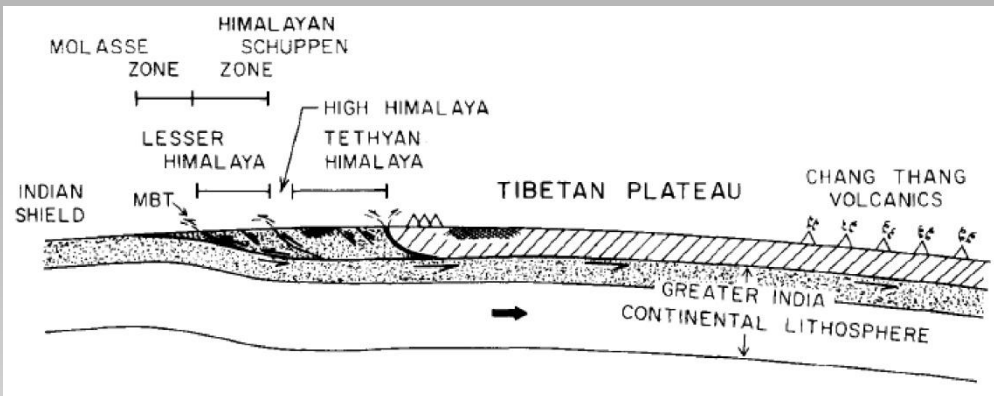
Wannamaker et al., (2002)



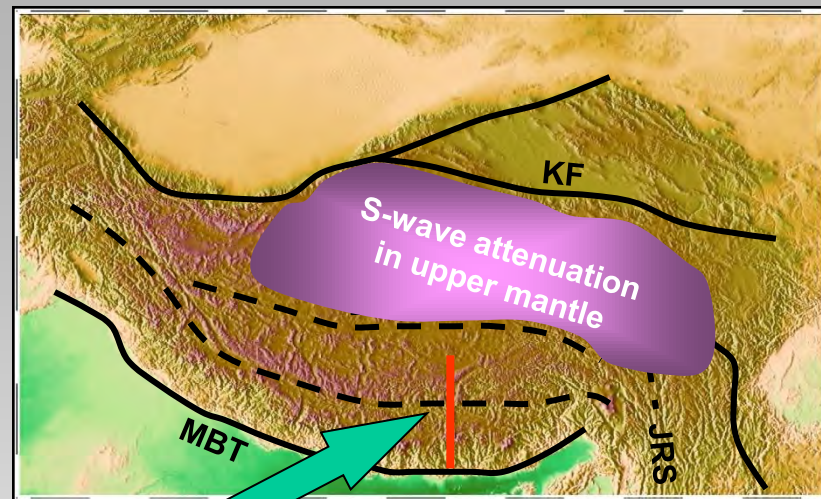
Bertrand et al., (2009)



13.7 : Fluids generated in collision zones (Himalaya and Tibetan Plateau)

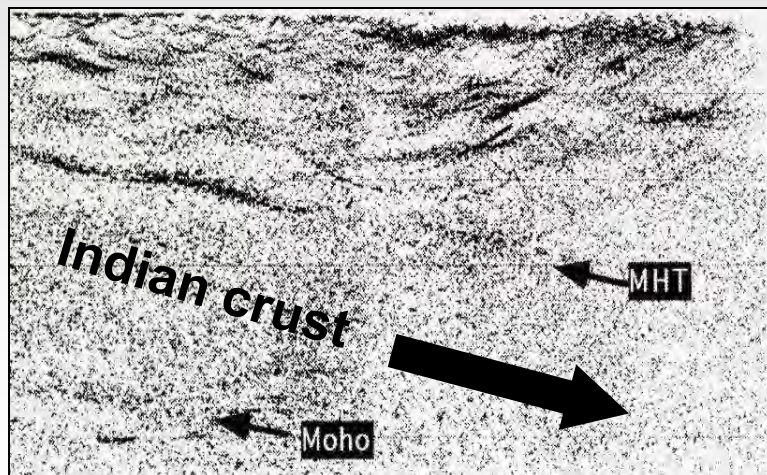


Ni and Baranzangi, *JGR*, (1982)



INDEPTH-I 1992

INDEPTH-II 1994



0 km

100 km



Nelson et al., *Science*, (1996)

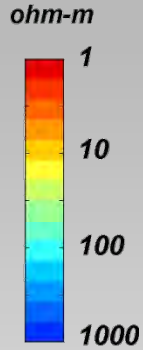
13.7 : Fluids generated in collision zones (Himalaya and Tibetan Plateau)

Magnetotelluric studies (1995 – 2001)



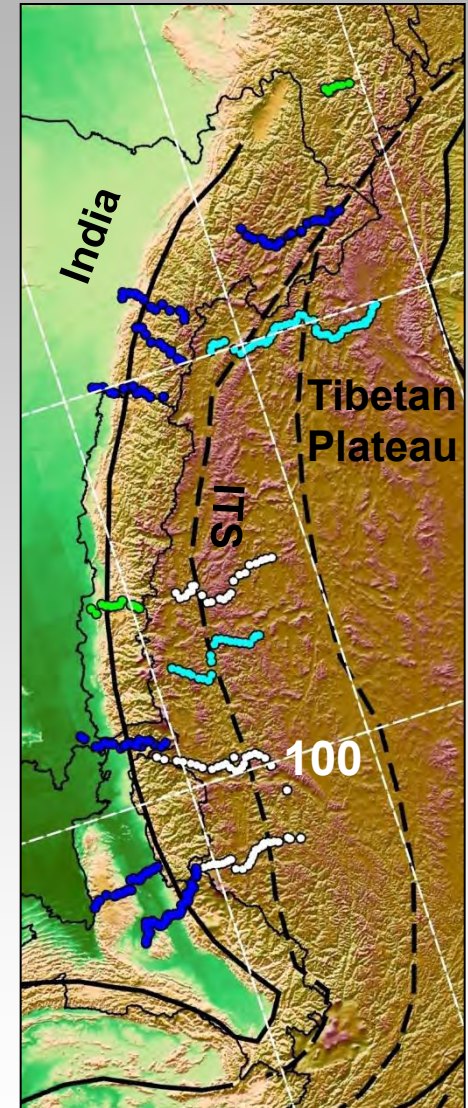
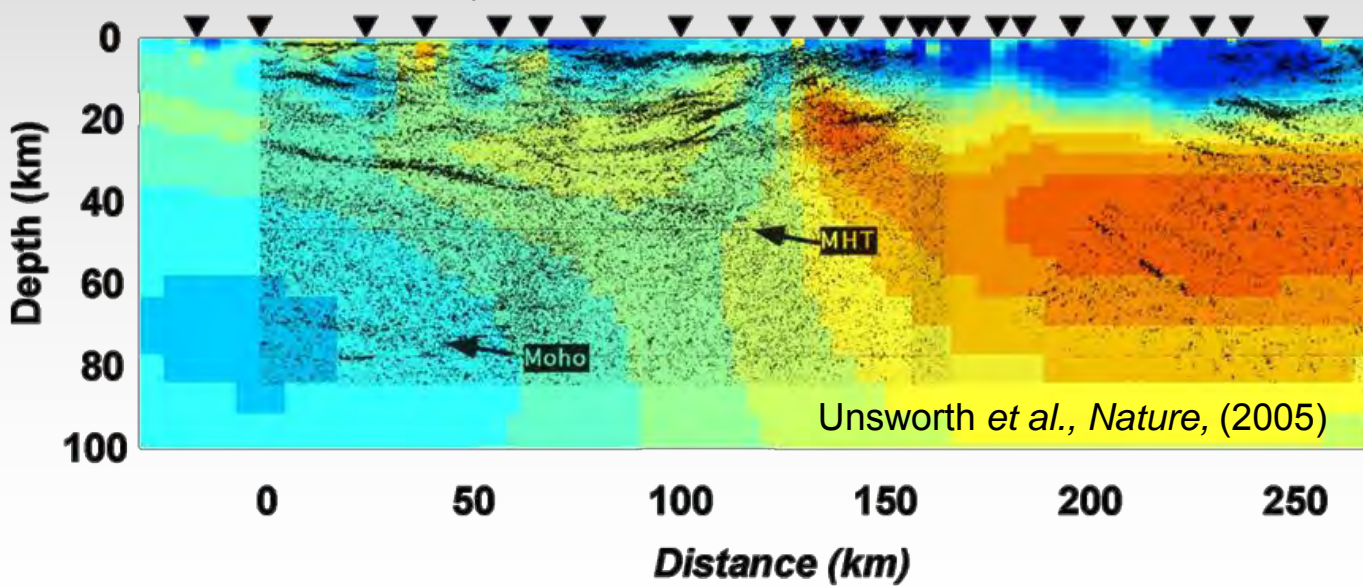
International **Deep Profiling** of **Tibet**
and the **Himalaya (INDEPTH)**

13.7 : Fluids generated in collision zones (Himalaya and Tibetan Plateau)

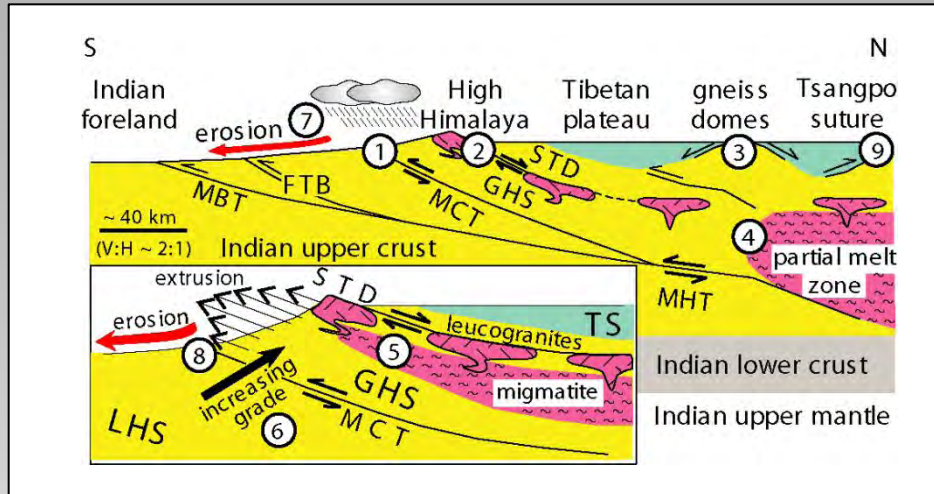


High
Himalaya

ITS

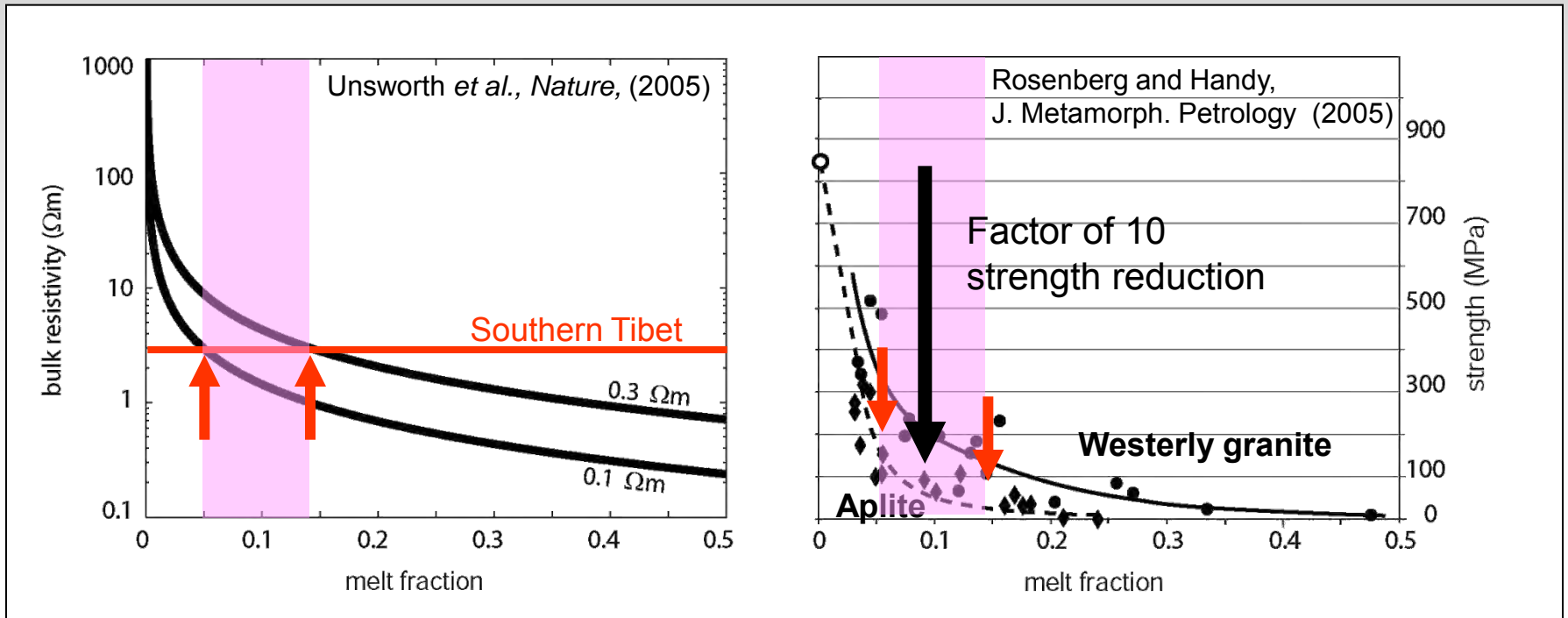


13.7 : Fluids generated in collision zones (Himalaya and Tibetan Plateau)

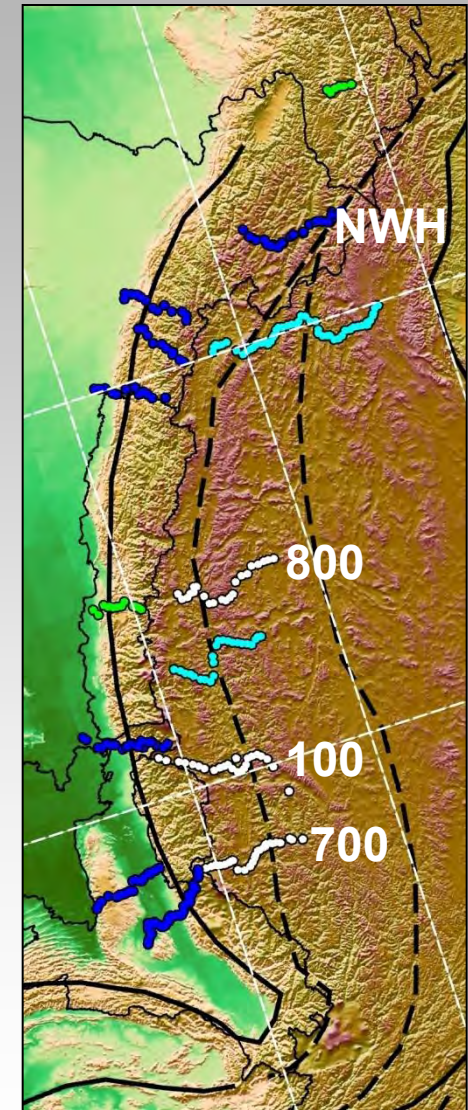
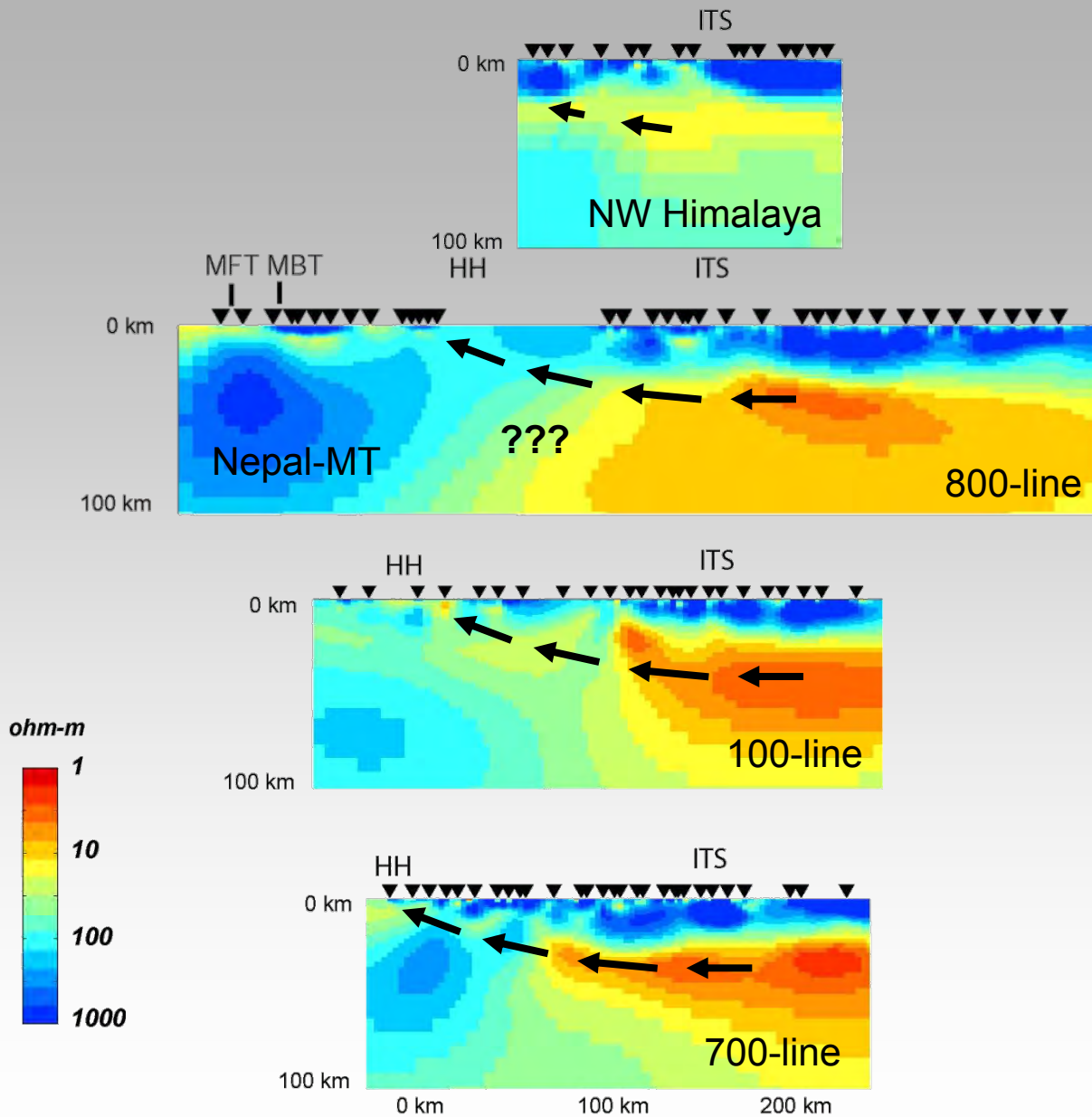


•Beaumont et al., *Nature* (2002) model requires factor of 10 reduction in viscosity for crustal flow to occur

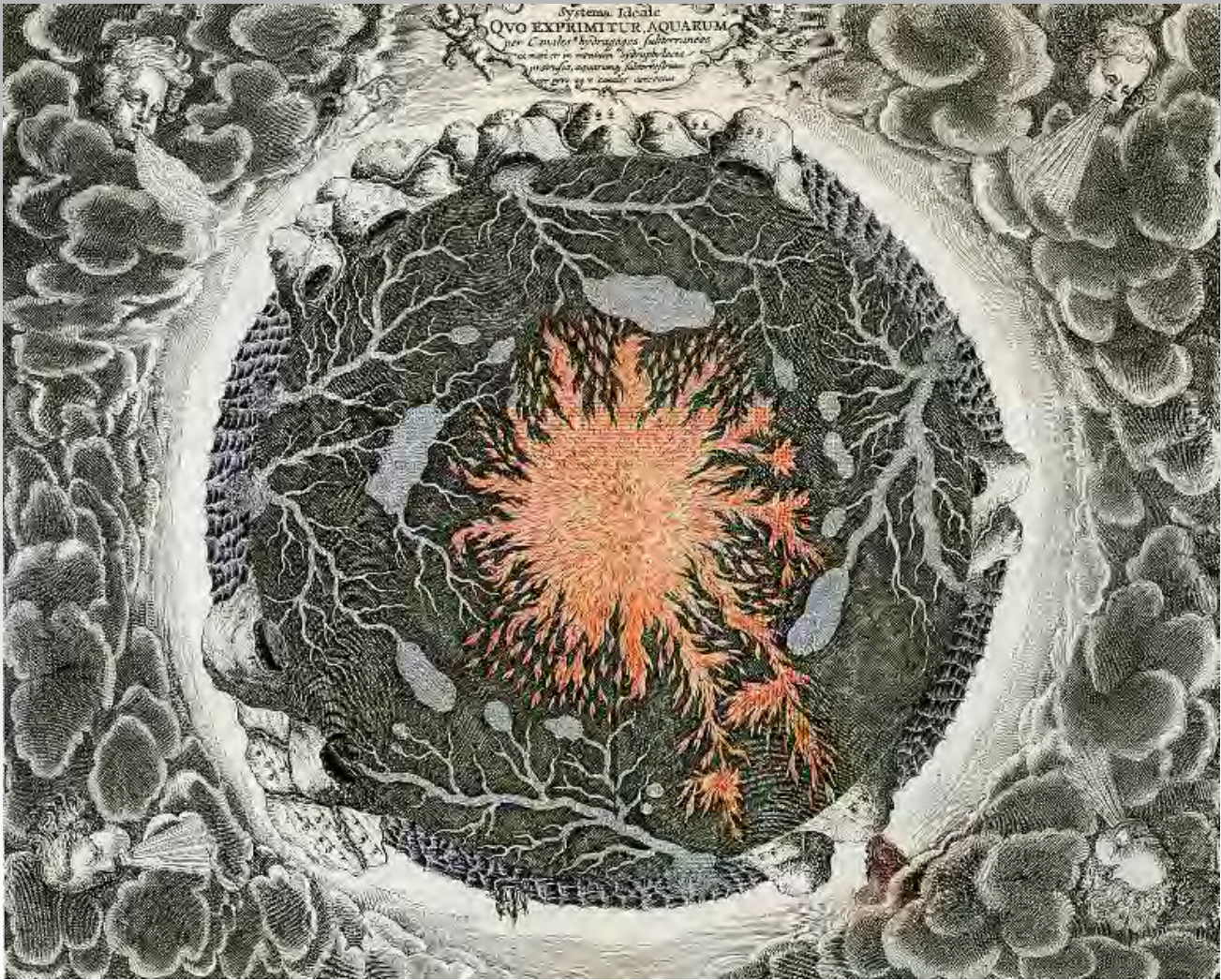
•MT data require 5-12% melt. This is consistent with a factor of 10 strength reduction.



13.7 : Fluids generated in collision zones (Himalaya and Tibetan Plateau)



13.8 : Conclusions



ACKNOWLEDGEMENTS

- Invitation from Dan Harlov and Tom Chacko to attend Goldschmidt meeting in 2008.
- Models and data from many colleagues.
- Funding from NSF, NSERC, DOE, US Geological Survey, Alberta Ingenuity fund and University of Alberta
- Many people who worked in the field to collect the data

